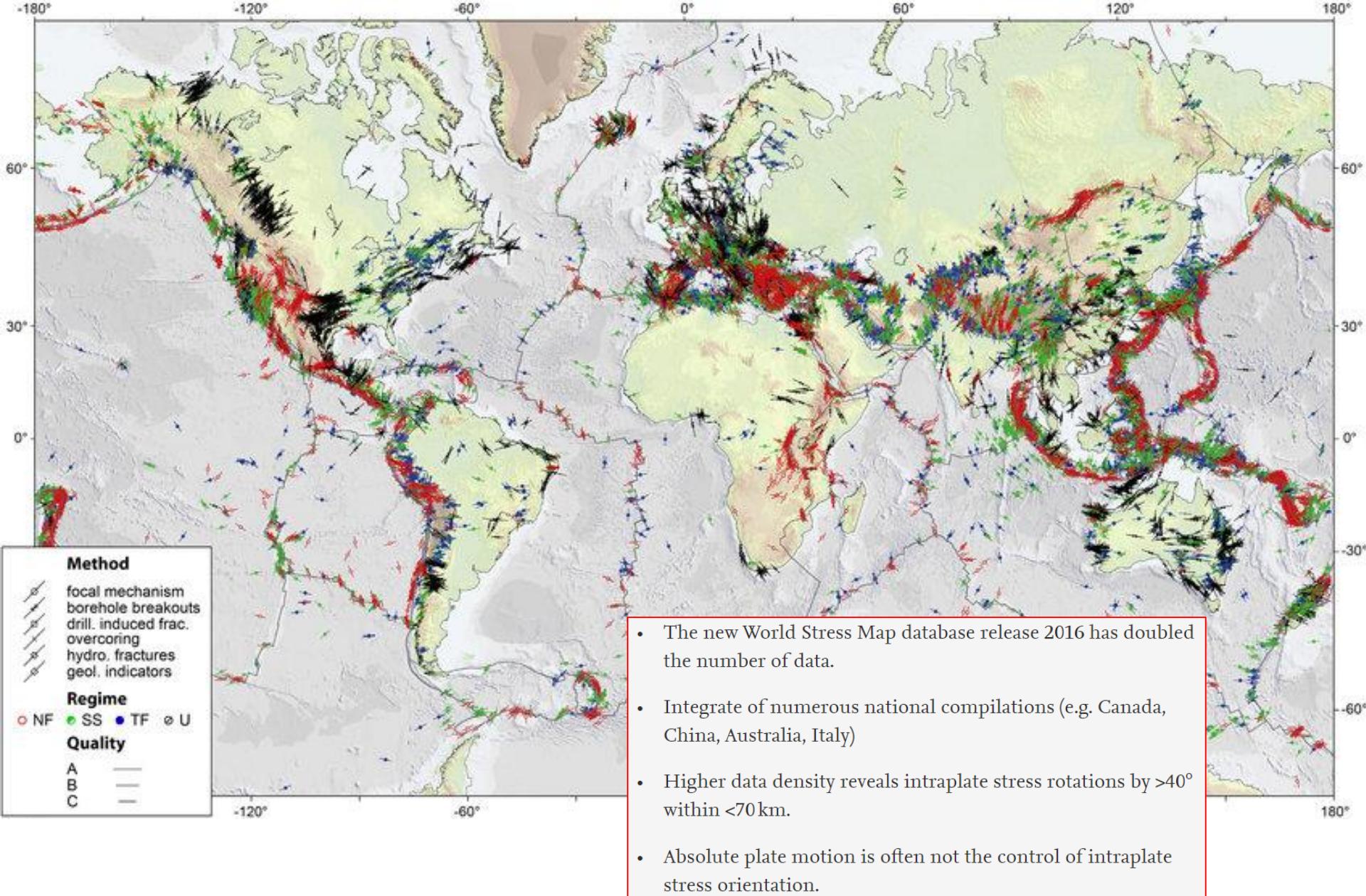


# ESTADO DE ESFUERZOS EN EL SUR Y NORTE DE HONDURAS

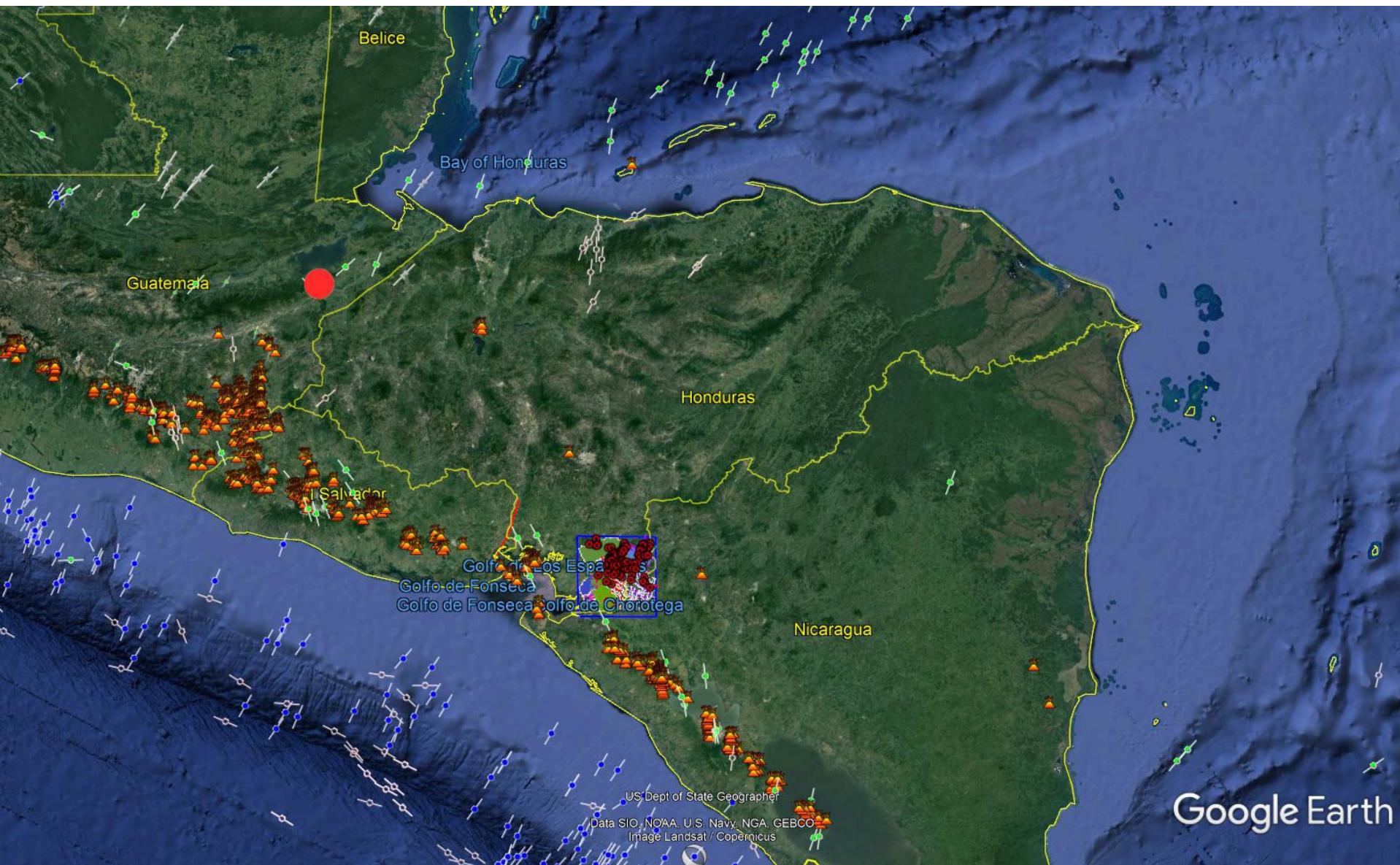
Allan López

*Centro de Investigaciones en Ciencias Geológicas  
Universidad de Costa Rica*

*Ingeniería Civil  
Universidad Latina de Costa Rica*



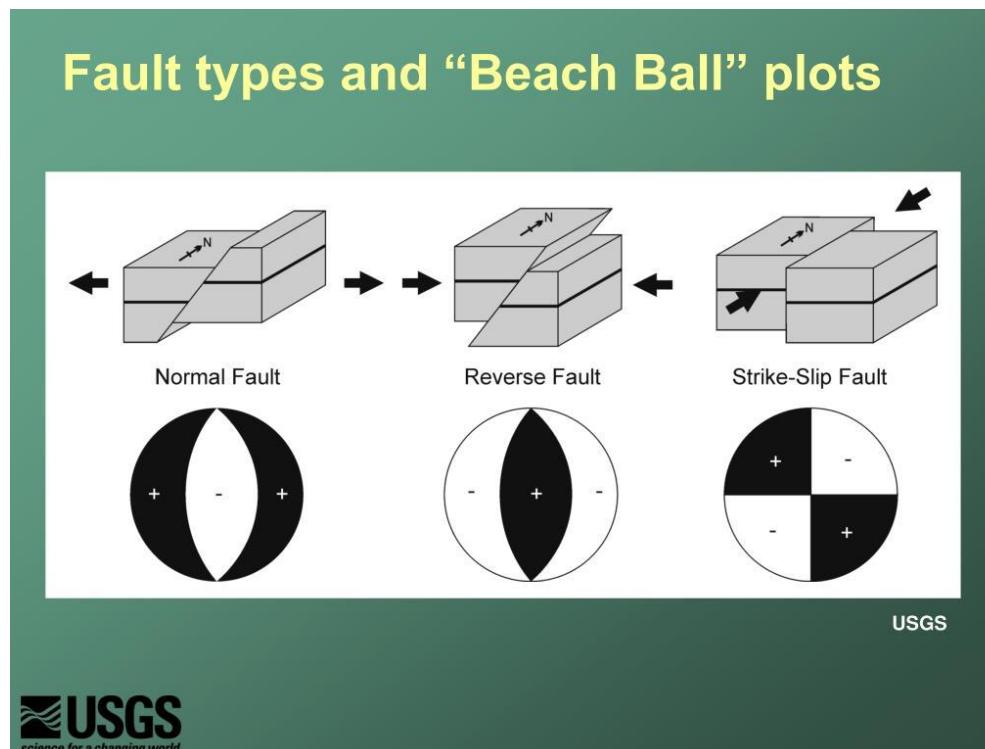
**The WSM database release 2016 contains 42,870 data records within the upper 40 km of the Earth's crust**

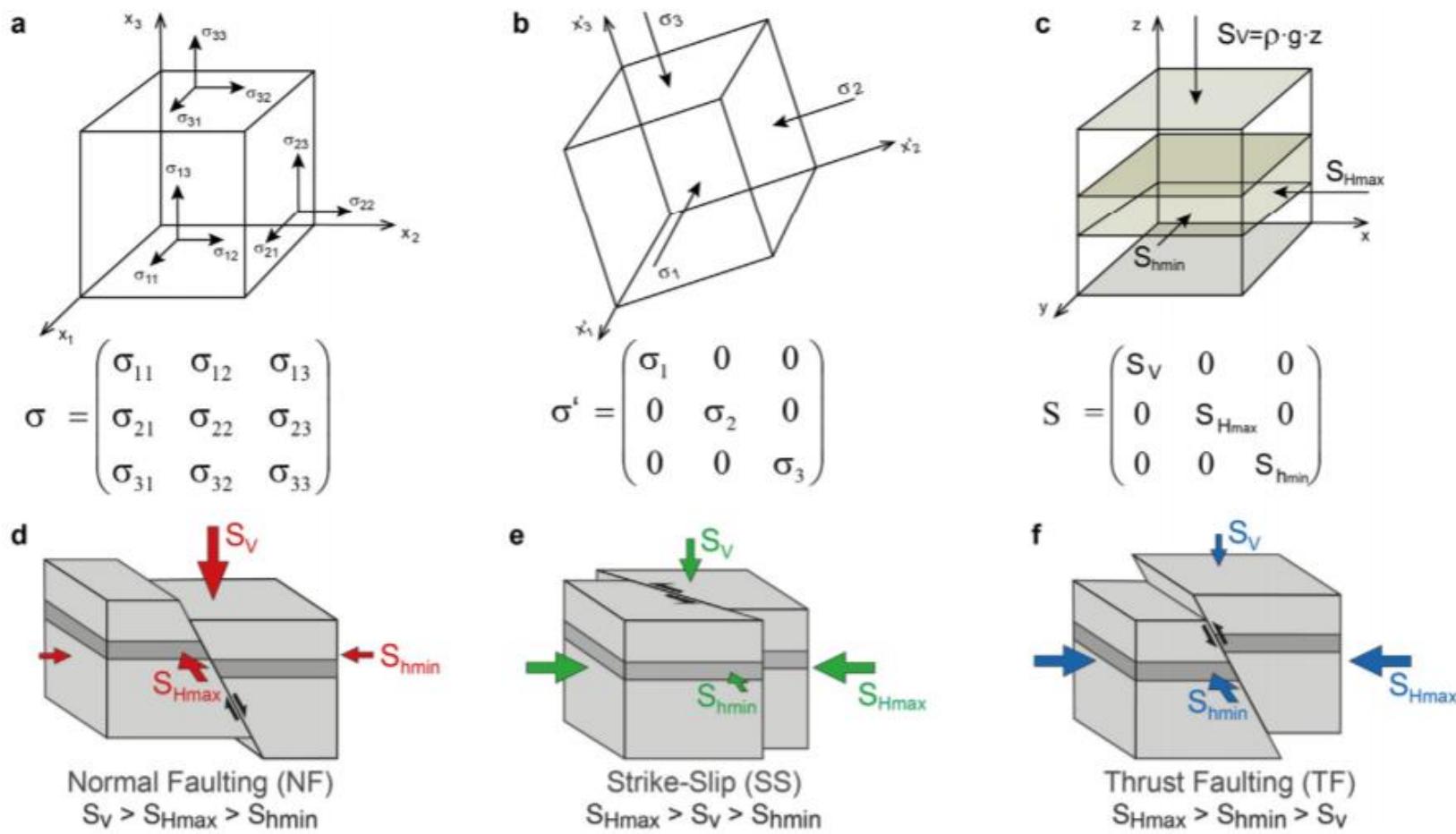


El conocimiento de las características y propiedades físicas y mecánicas del **estado de esfuerzo tectónico** prevaleciente en la corteza terrestre es de fundamental importancia para la sociedad y la industria, ya que tiene implicaciones y aplicaciones directas en:

- La caracterización de reservorios geotérmicos, de petróleo, gas y su administración racional y segura.
- La estabilidad de minas, túneles carreteros y de conducción de agua, vertederos de desechos tóxicos y no tóxicos.  
En general para las obras subterráneas.
- Calibración de modelos numéricos geo-mecánicos.
- Simulaciones termo, hidromecánicas en 4D.
- Evaluación del peligro y amenaza sísmica, mediante el análisis de la **Tendencia al deslizamiento de la falla y el potencial de fractura.**

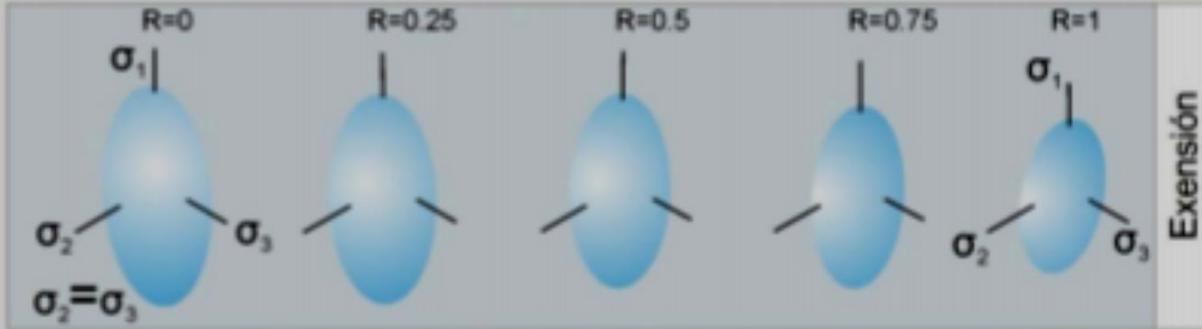
- Mecanismos focales de terremotos
- Elongaciones-breakouts en pozos y fracturas inducidas por la perforación
- Medidas In-situ (overcoreing, fracturación hidráulica, borehole slotter)
- Datos Geológicos recientes (análisis de fallas mesoscópicas y alineamiento de conos y diques volcánicos Cuaternarios)



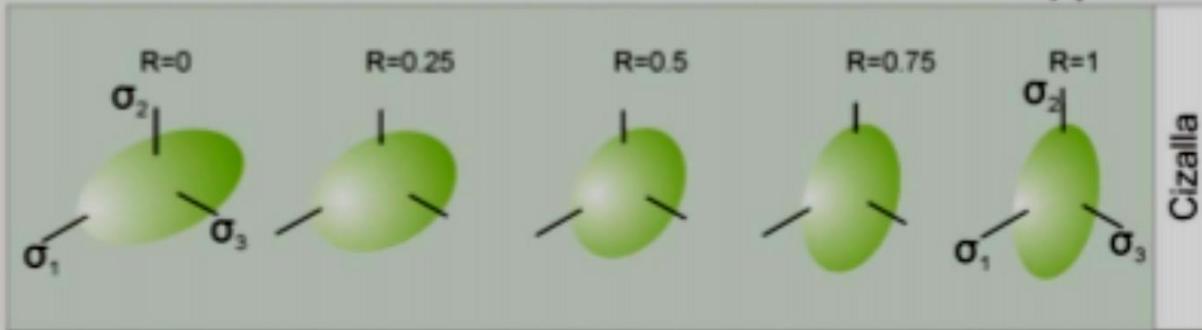


**Fig. 1.** a) The components of the stress tensor define the stress state at a point and enable to compute the stress vector on any surface within a body. To describe the stress tensor components an infinitely small cube with unit surfaces is used. The forces acting on the cube faces can be decomposed in forces acting parallel and those acting normal to the surface. The first create the normal stress components, the second create the shear stress components of the stress tensor. b) Due to the conservation of momentum the stress tensor has to be symmetric. This implies that a coordinate system exists where shear stresses vanish along the faces of the cube. In this principal axis system the remaining three stresses are the principal stresses. c) Assuming that the vertical stress in the Earth crust  $S_V = g \cdot \rho \cdot z$  is a principal stress ( $g$  is gravitational acceleration,  $\rho$  is the rock density,  $z$  the depth), the two horizontal stresses  $S_{h\min}$  and  $S_{H\max}$ , the minimum and maximum horizontal stress, respectively, are principal stresses as well. This so-called reduced stress tensor is fully determined with four components: the  $S_{H\max}$  orientation and the magnitudes of  $S_V$ ,  $S_{H\max}$  and  $S_{h\min}$ . The stress information provided in the WSM database is for all data records the  $S_{H\max}$  orientation and for most data records the stress regime which indicates the relative stress magnitudes. d) Normal faulting stress regime where  $S_V$  is larger than the horizontal stresses ( $S_V = \sigma_1$ ). e) Strike-slip faulting stress regime where  $S_V$  is the intermediate principal stress ( $S_V = \sigma_2$ ). f) Thrust faulting stress regime where  $S_V$  is smaller than the horizontal stresses ( $S_V = \sigma_3$ ).

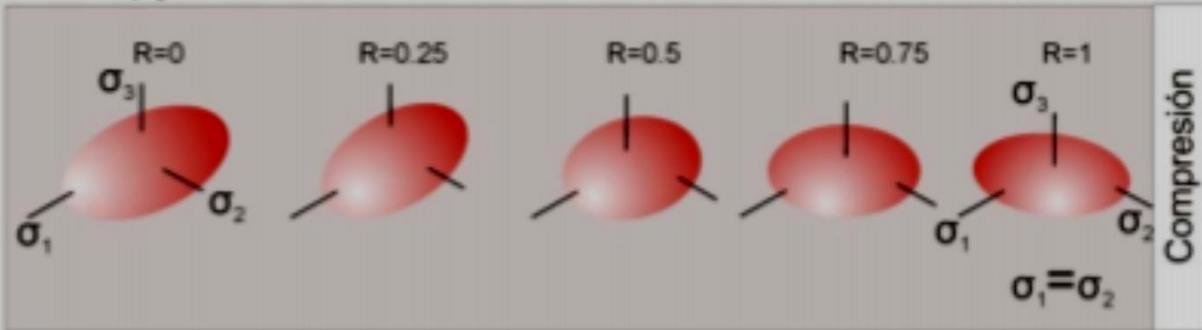
Radial ← Triaxial → Uniaxial



$$|| \sigma_1 = \sigma_2$$

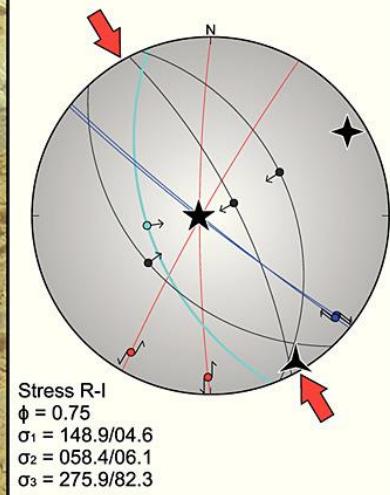
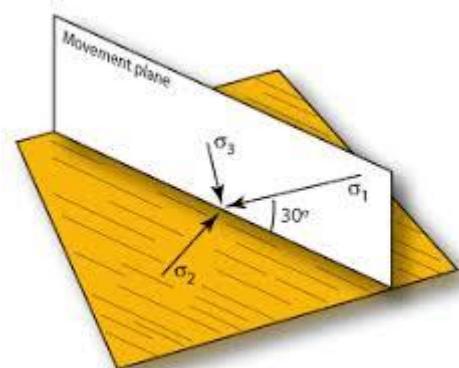
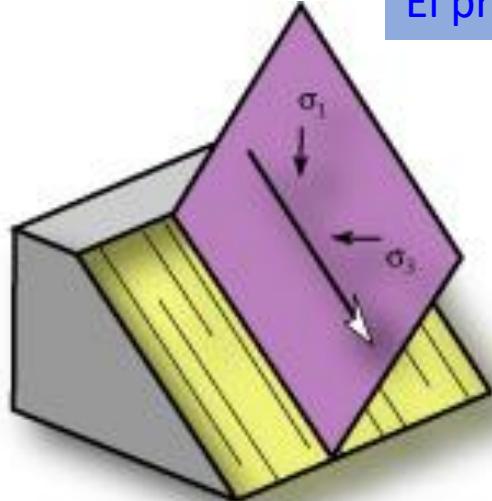


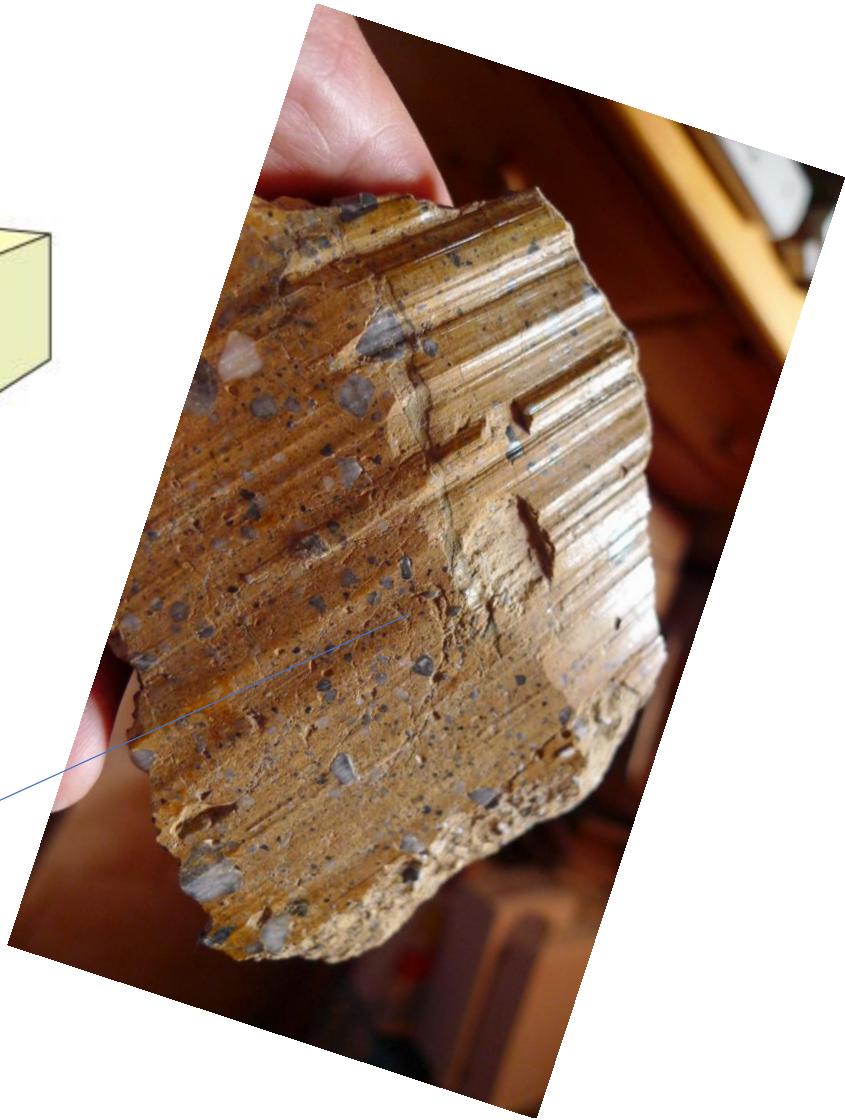
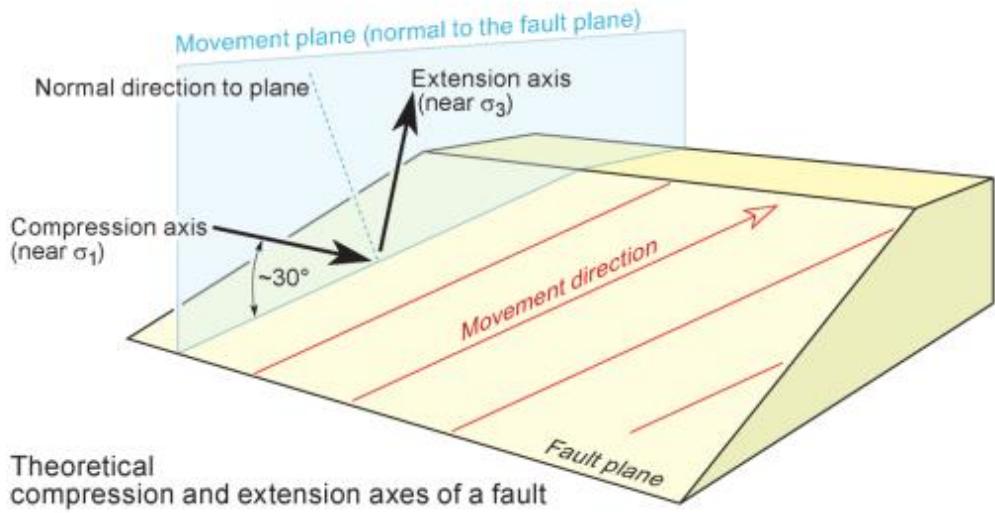
$$\sigma_2 = \sigma_3 ||$$

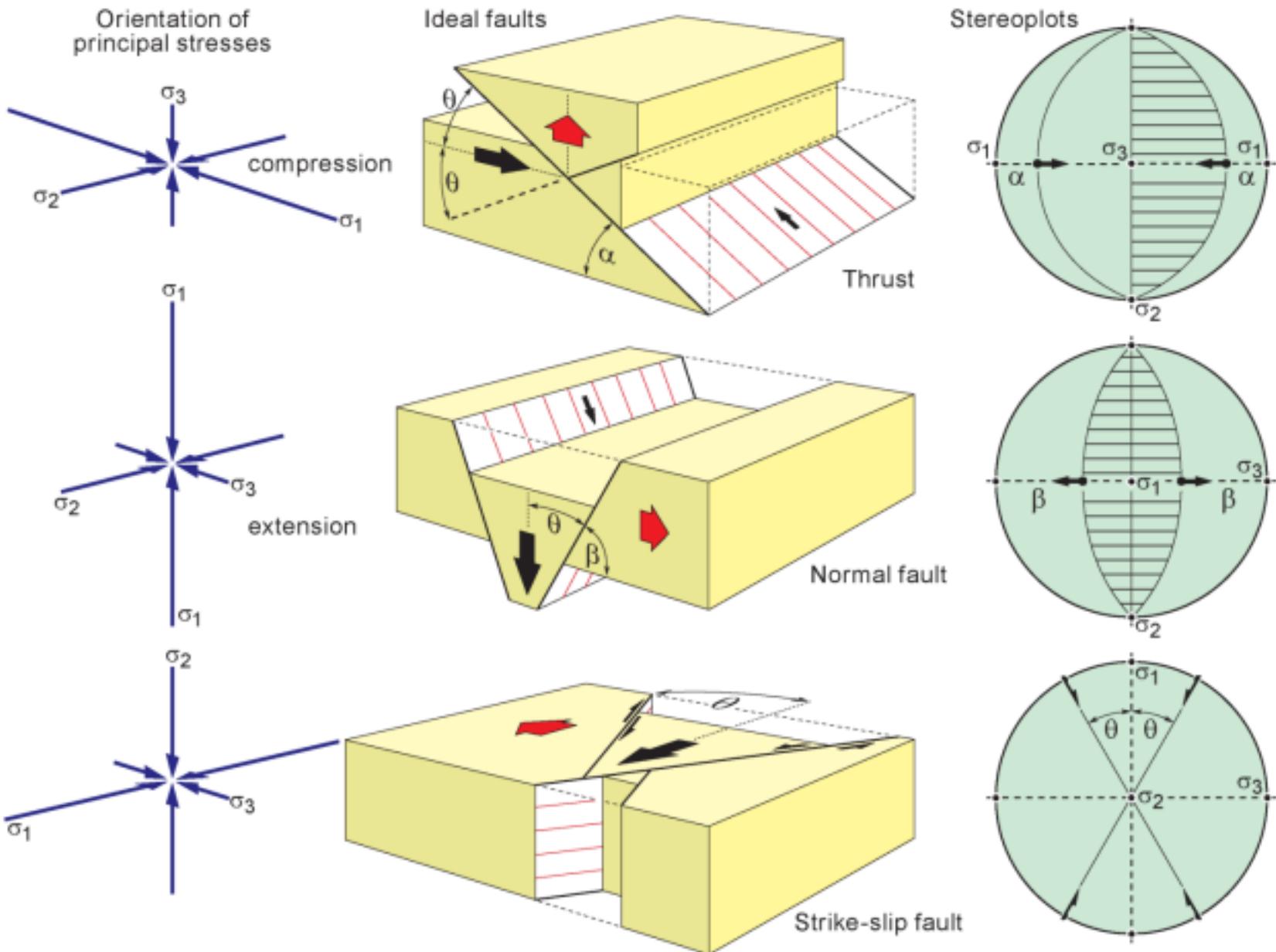


Uniaxial ← Triaxial → Radial

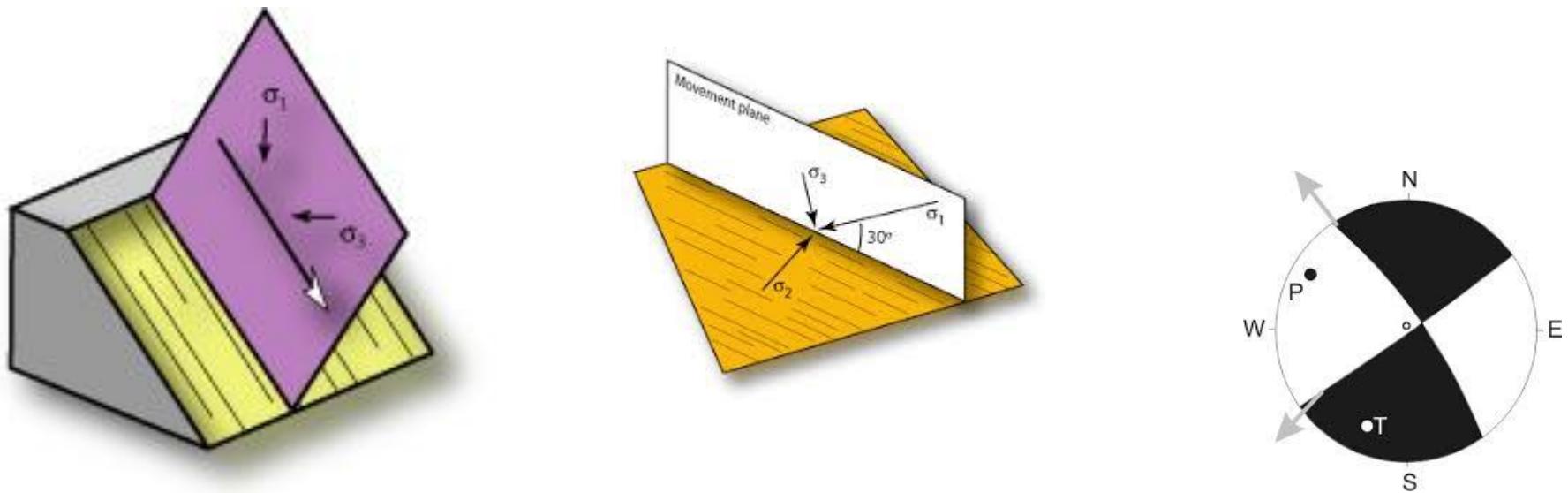
## El problema inverso en los esfuerzos tectónicos





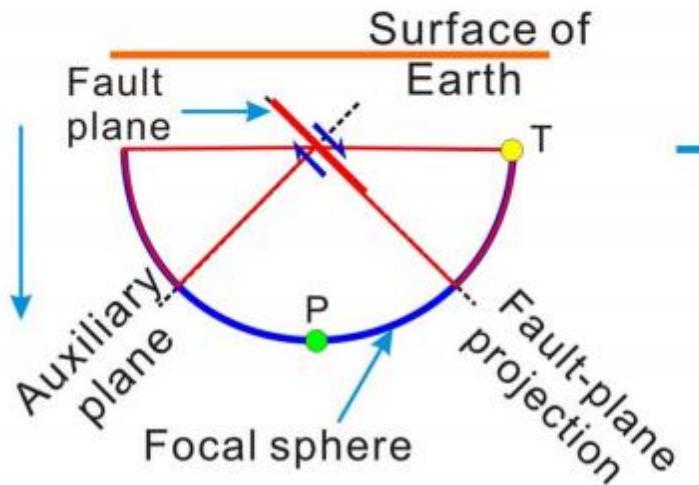


Dynamic interpretation of faults: Anderson's "standard" relationship between stresses and ideal faults

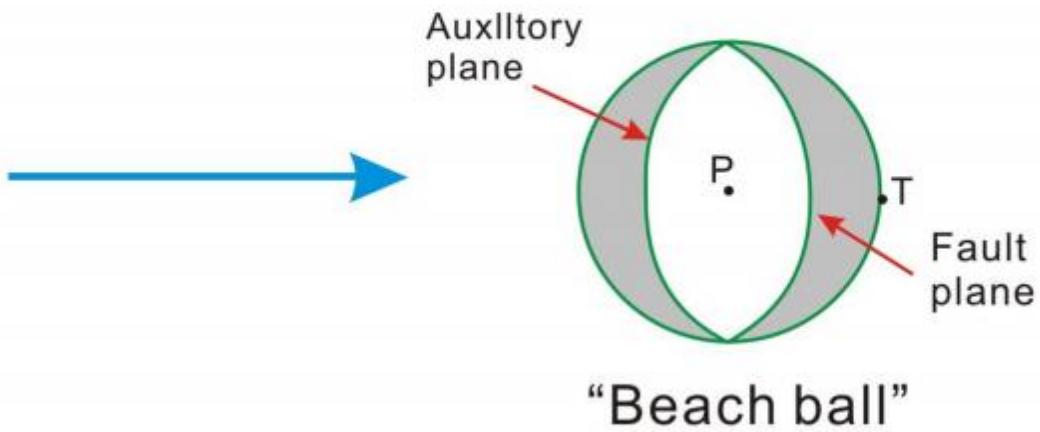


## Schematic diagram of a focal mechanism

View from side



View from above



$$\Phi = \frac{\sigma_2 - \sigma_3}{\sigma_1 - \sigma_3}$$

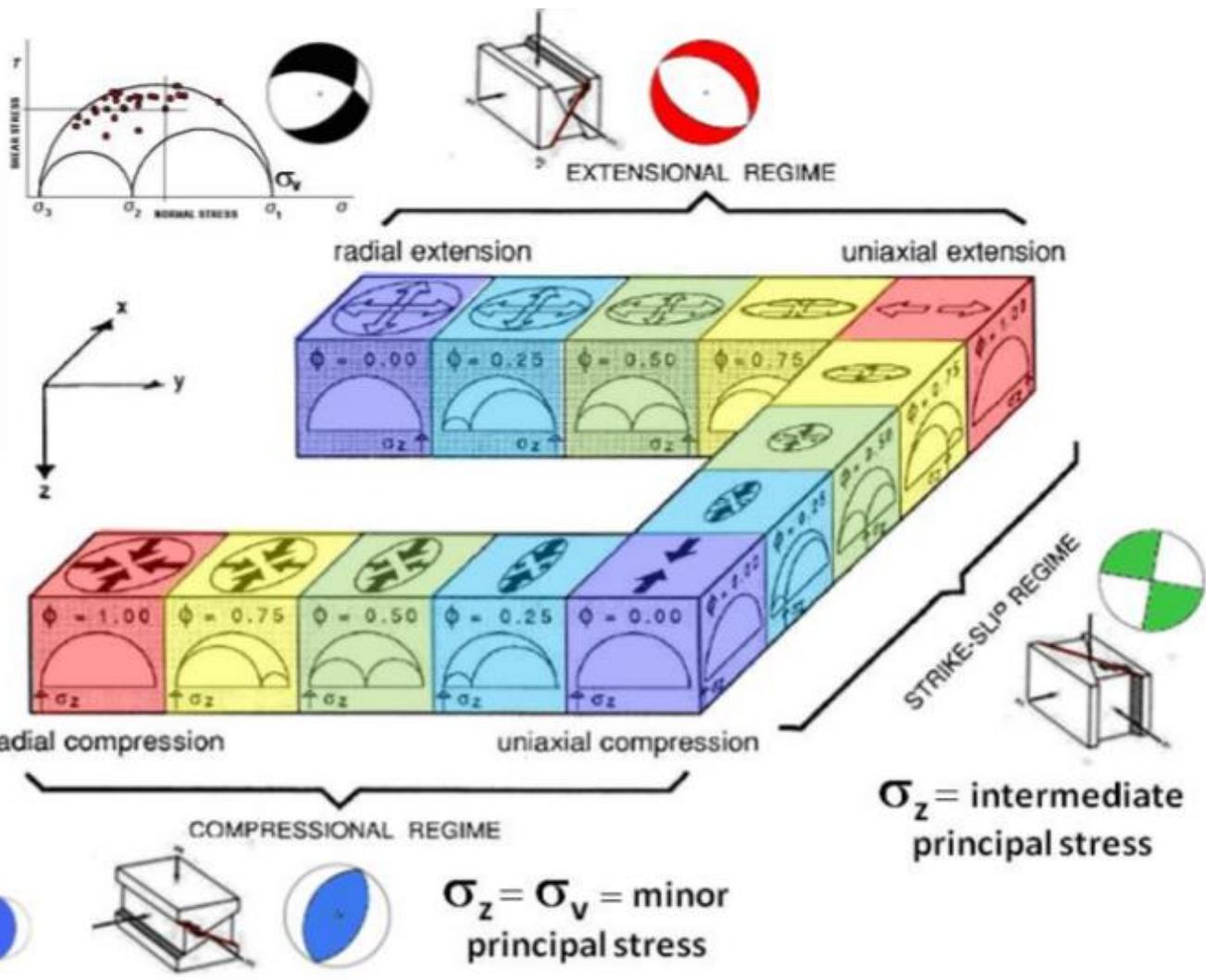
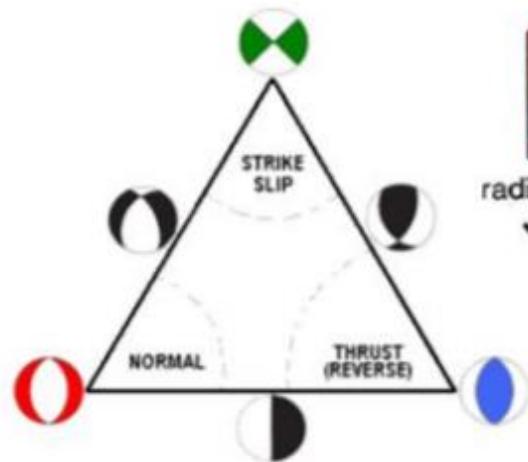
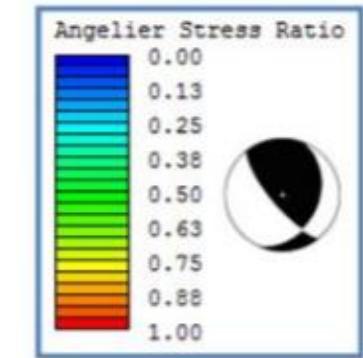
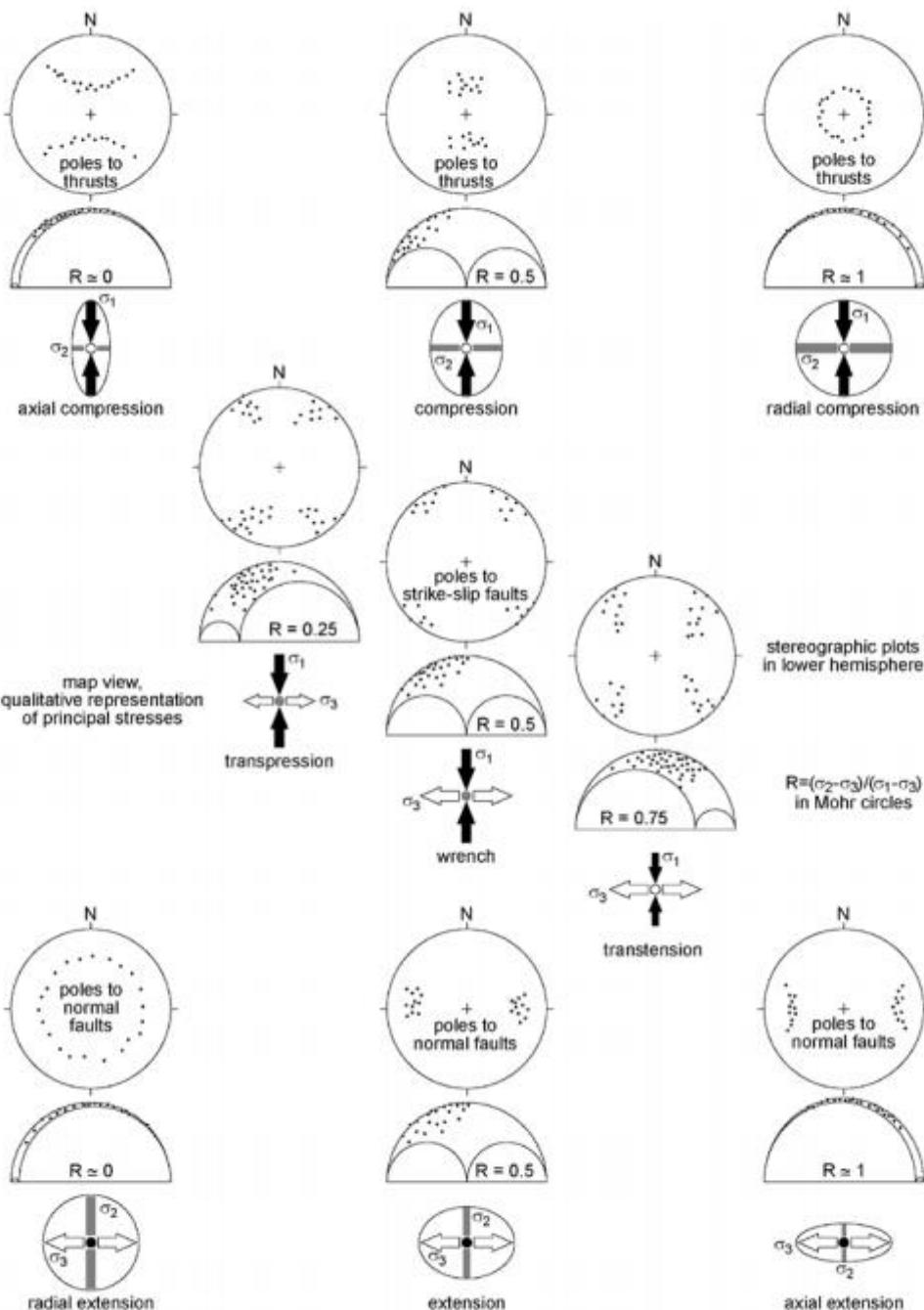


Figure 9 Classification of different tectonic stress states for different Fault geometries as plotted from beachballs and ternary diagrams, after Carter (2015b) (with key inset diagram modified from original created by Ritz 1994)

Por eso, NO hay que decir esta falla es normal, inversa, de rumbo, sino . . .  
el último movimiento conocido es . . . ( y en el futuro puede reactivarse con otra cinemática)



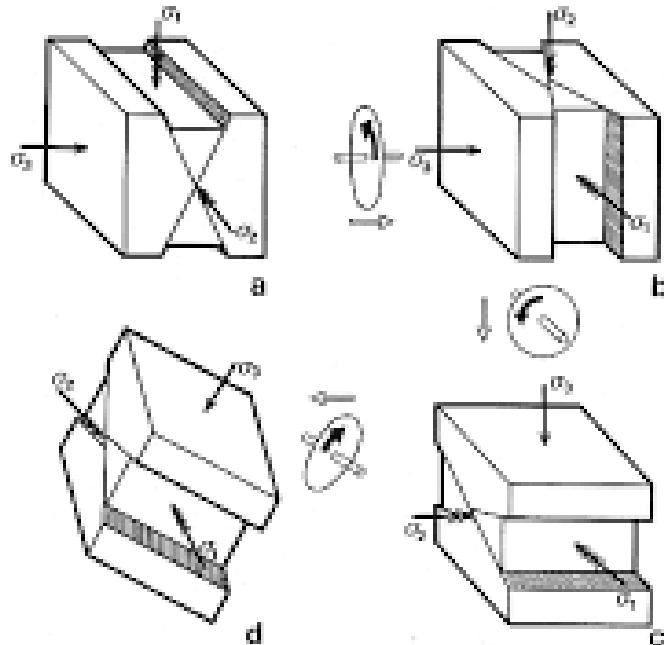
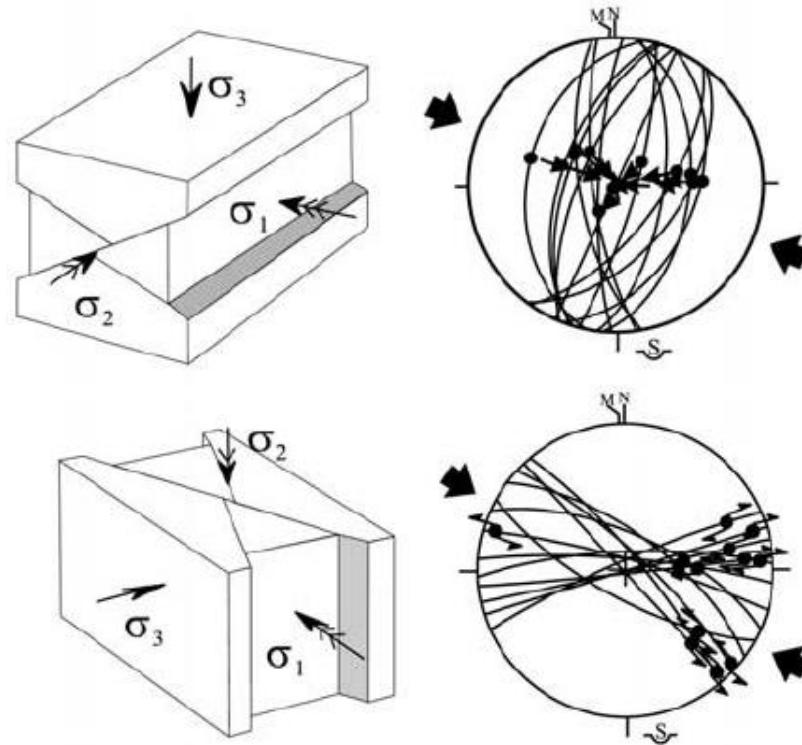
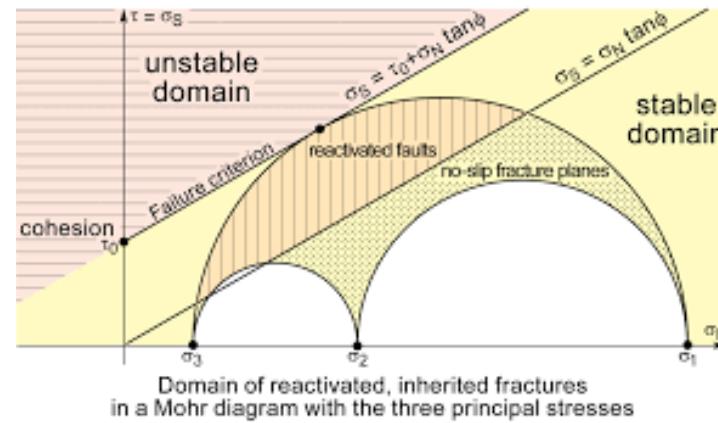
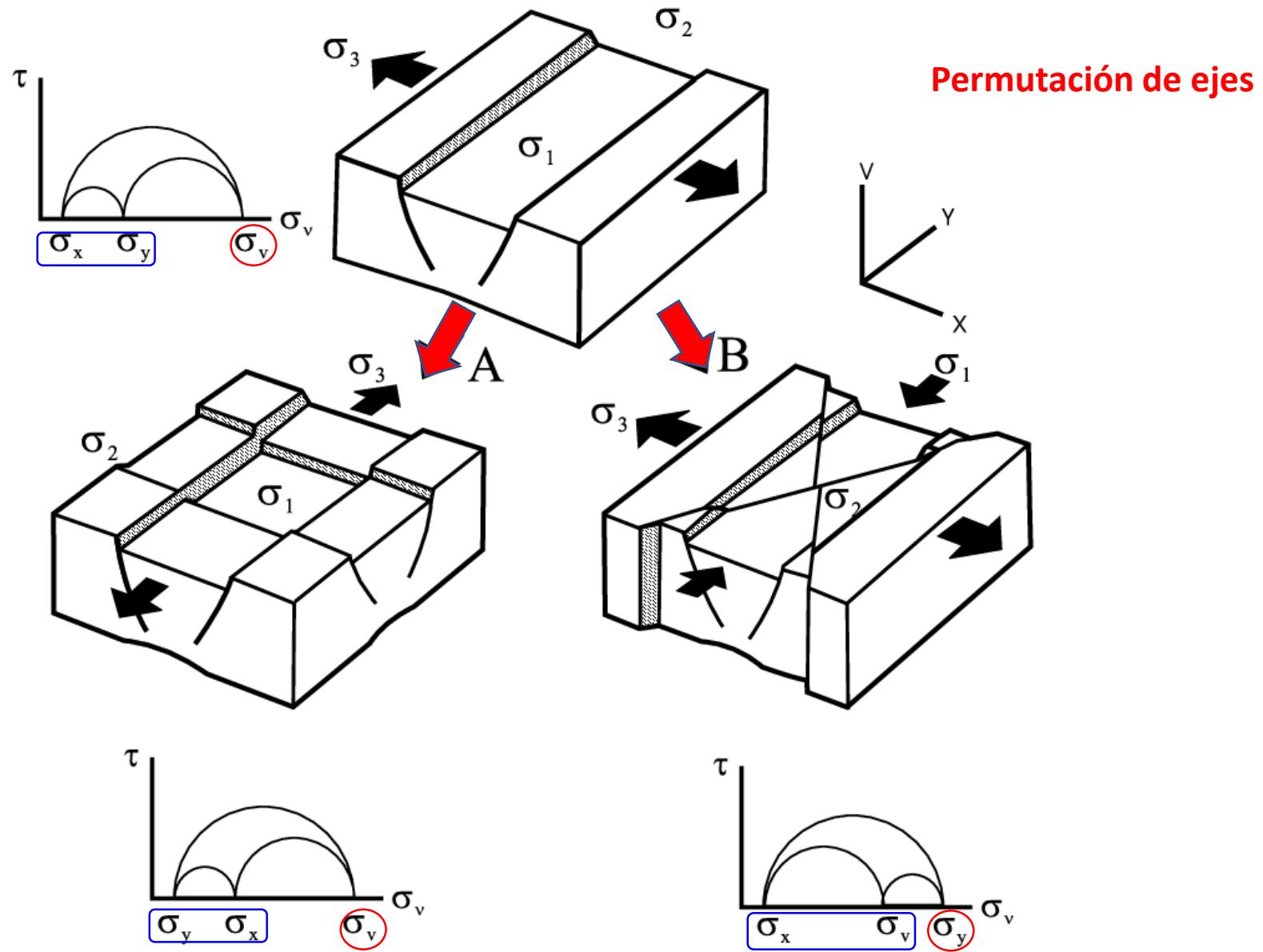


Fig. 4.18. Conjugate fault patterns. (a) Conjugate normal dip-slip faults. (b) Conjugate strike-slip faults. (c) Conjugate reverse dip-slip faults. (d) Conjugate oblique-slip faults. Principal stress axes shown with triple, double and single arrow-heads ( $\sigma_1$ ,  $\sigma_2$  and  $\sigma_3$ , respectively). Wheels with black arrows indicate rotations necessary to obtain (b) from (a), (c) from (b), and (d) from (c). Each of the first two rotations equals  $90^\circ$ .

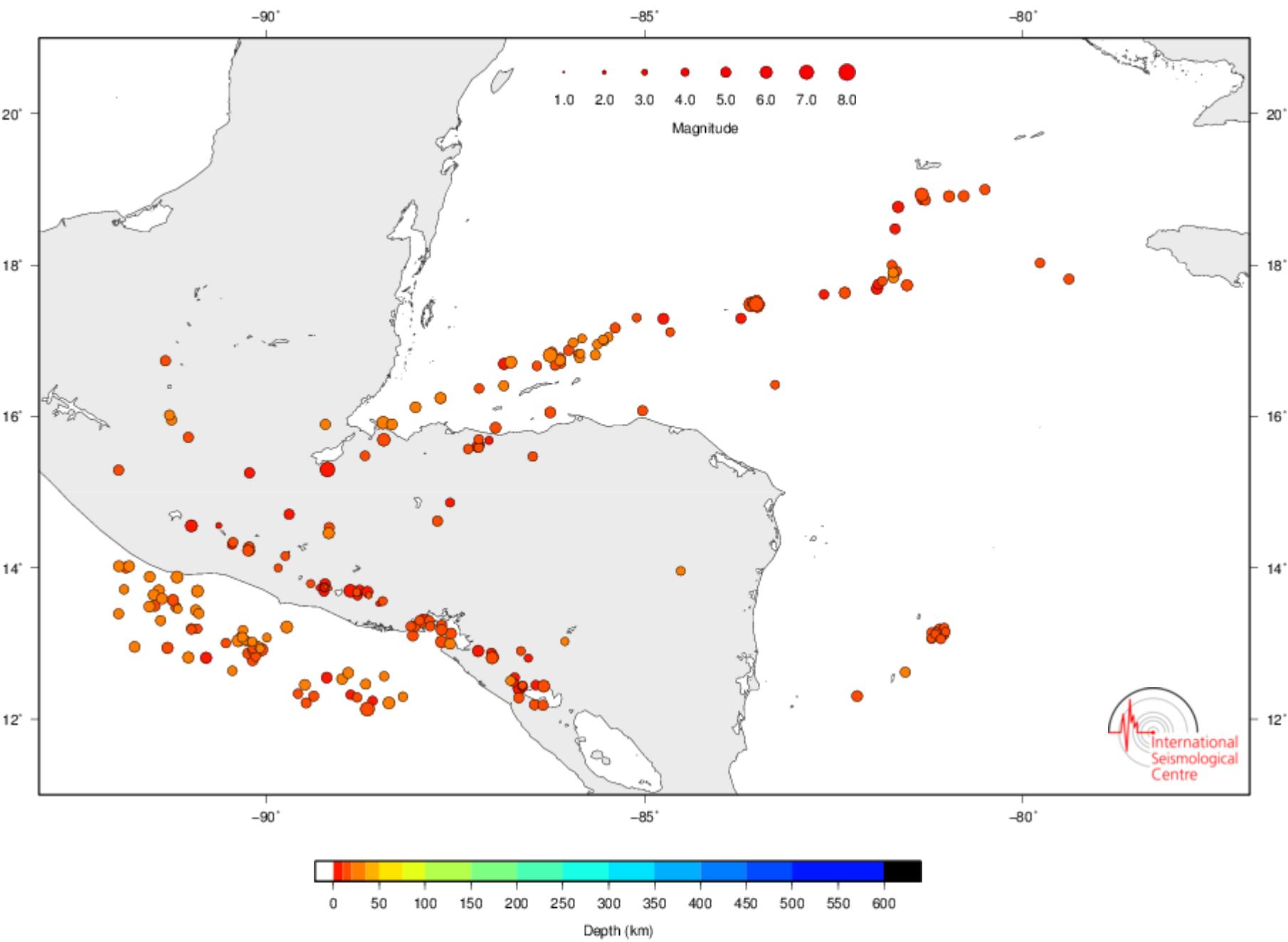


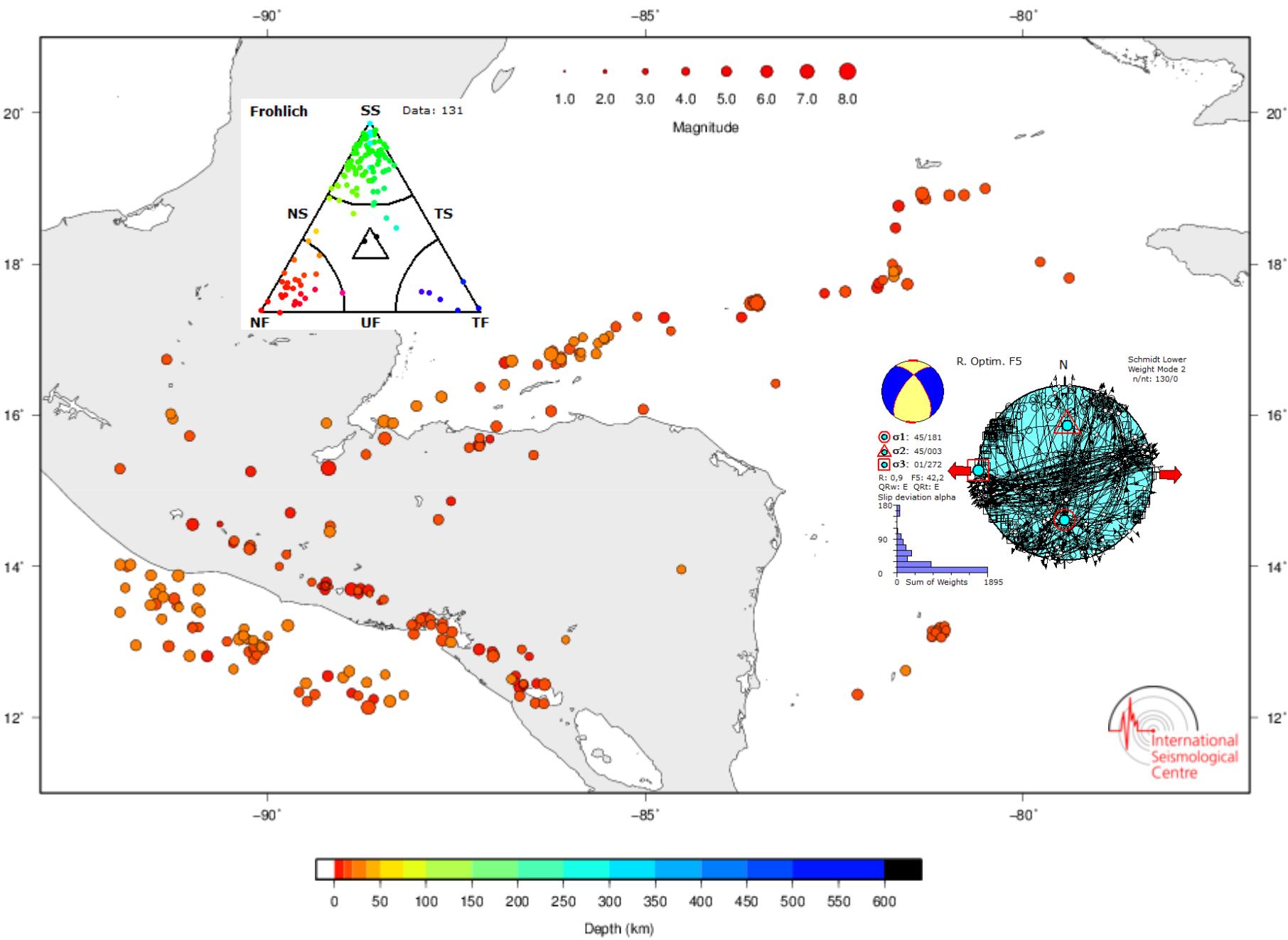
Por eso, NO hay que decir esta falla es normal, inversa, de rumbo, sino . . .  
el último movimiento conocido es .... ( y en el futuro puede reactivarse con otra cinemática)



Por eso, NO hay que decir esta falla es normal, inversa, de rumbo, sino . . .  
el último movimiento conocido es .... ( y en el futuro puede reactivarse con otra cinemática)

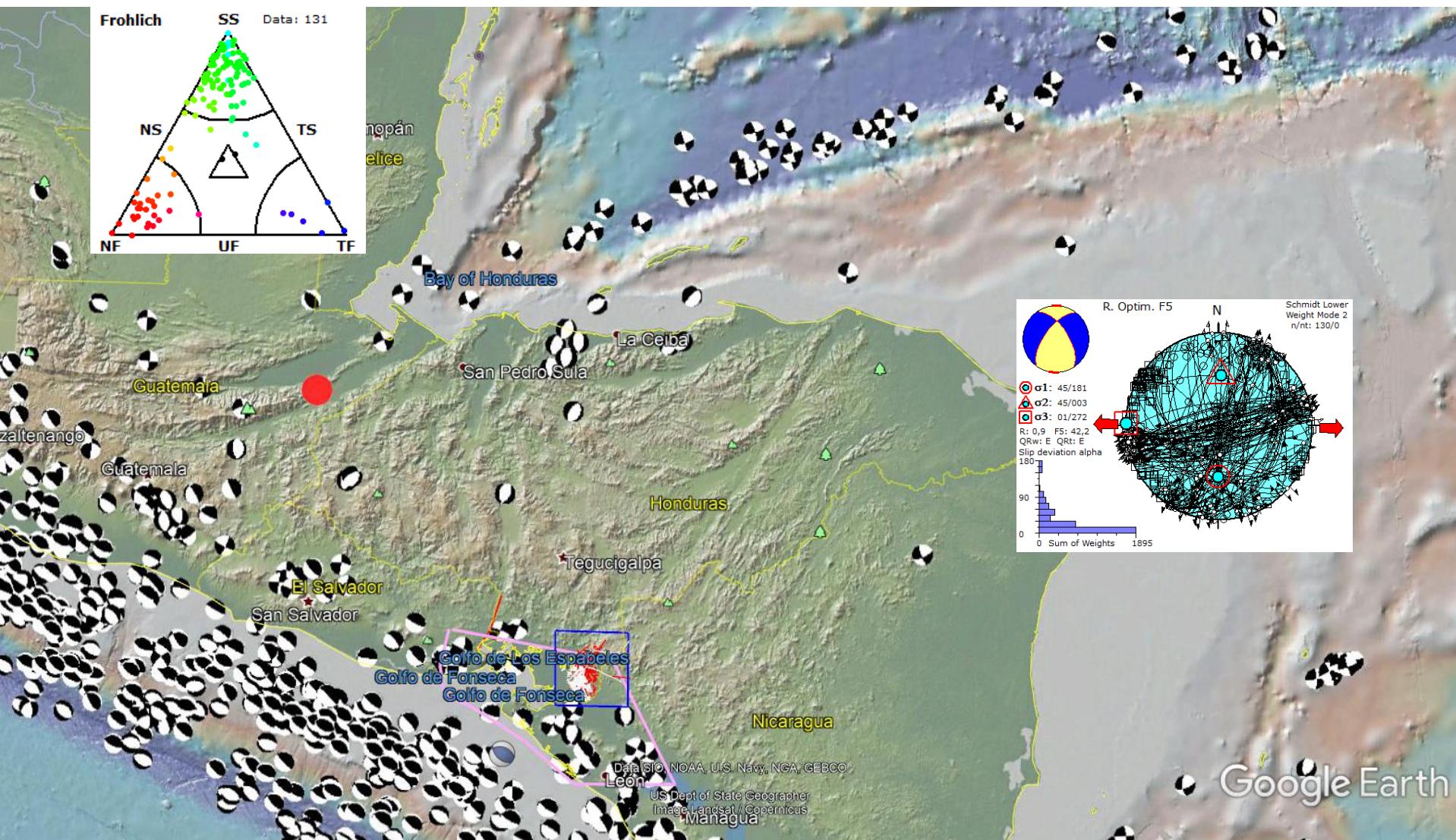
# ESTADO DE ESFUERZOS EN EL SUR Y NORTE DE HONDURAS



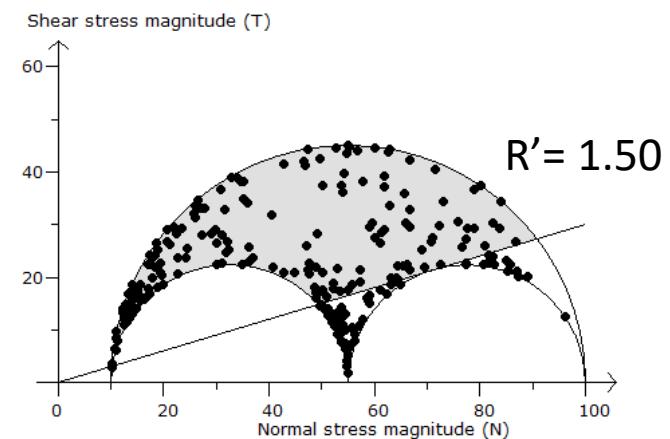
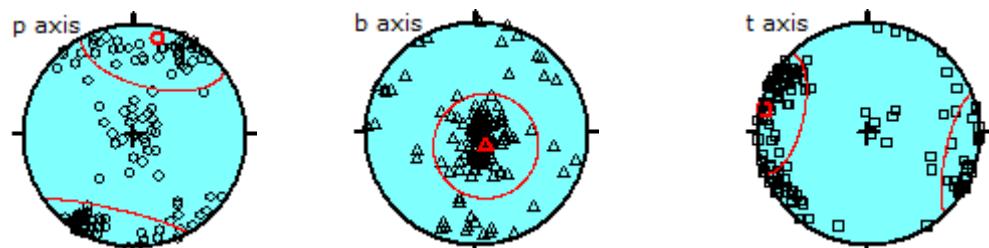
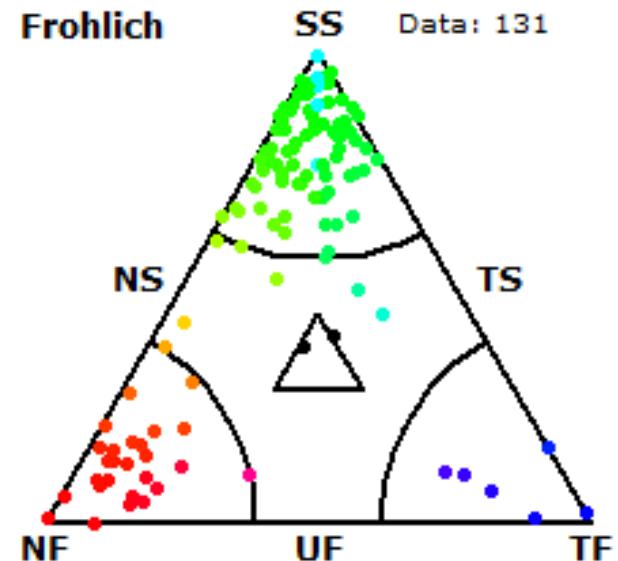
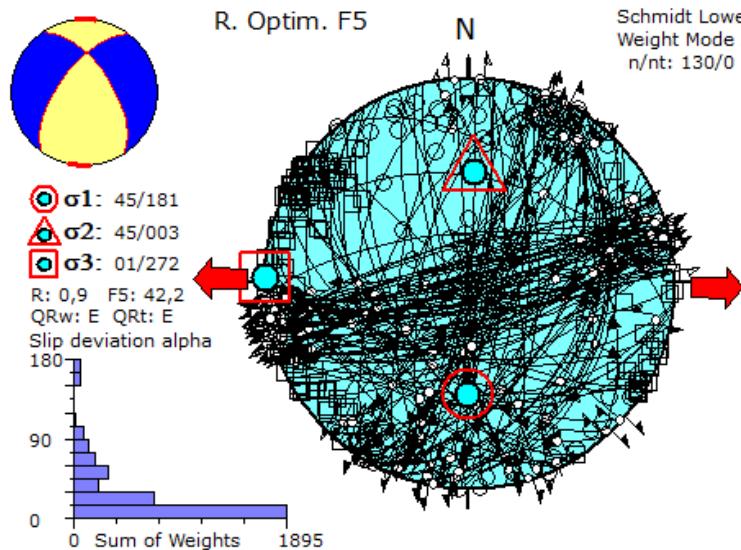


# Mecanismos focales regionales CMT

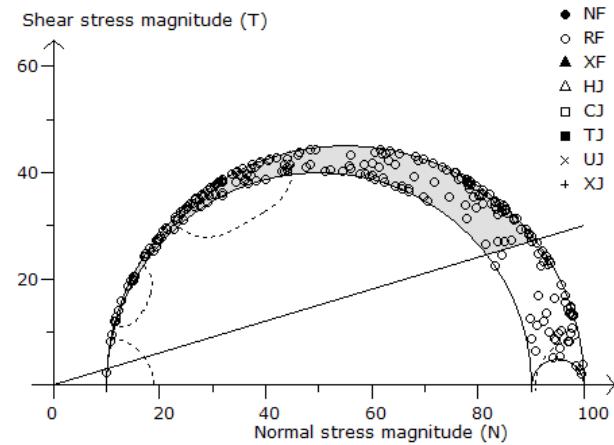
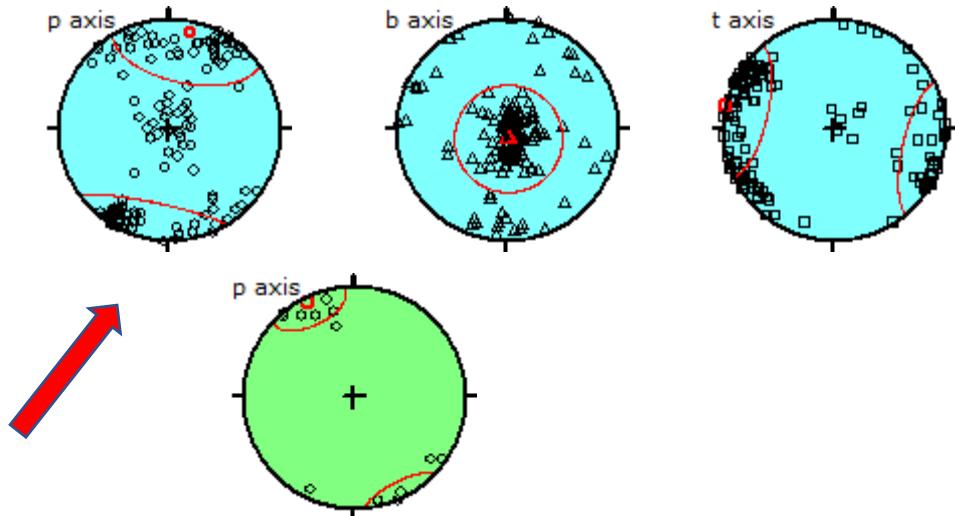
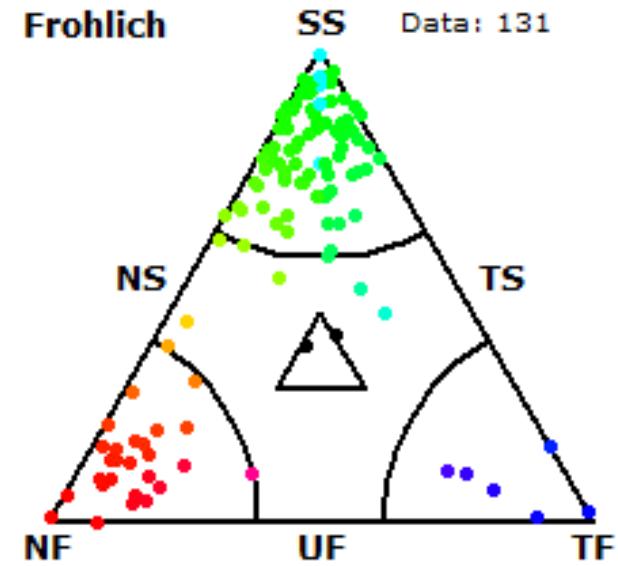
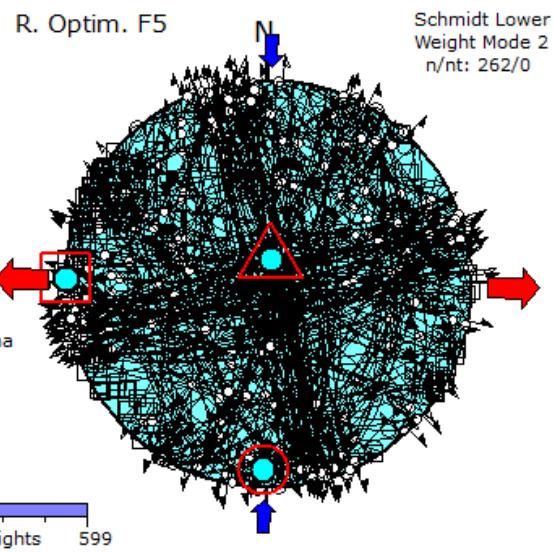
## Escena en GeoMapApp



# 1 Plano nodal de cada evento Mw > 4.0

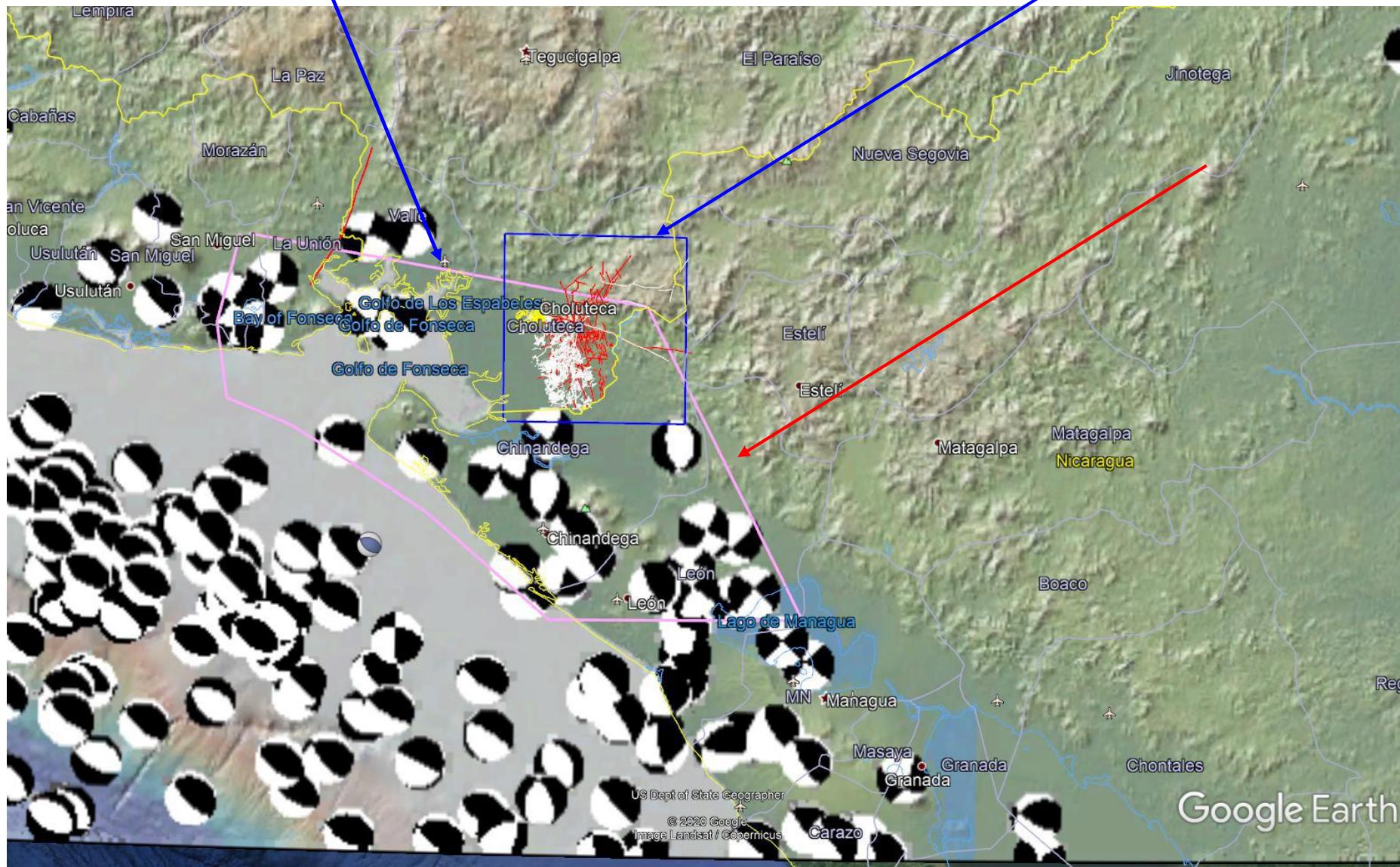


## 2 Planos nodales de cada evento Mw > 4.0



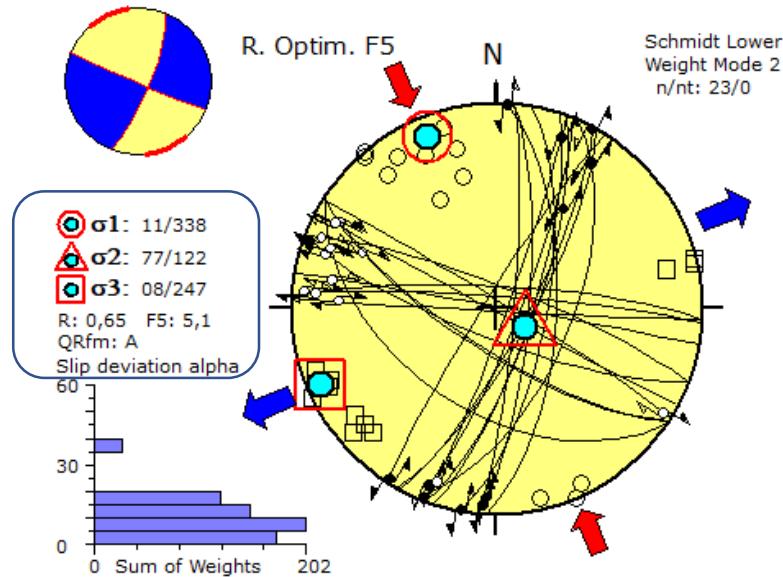
**Región con mecanismos focales  
seleccionados para calcular  
estado de esfuerzos tectónicos LOCAL**

**Región MED SRTM 30 m**

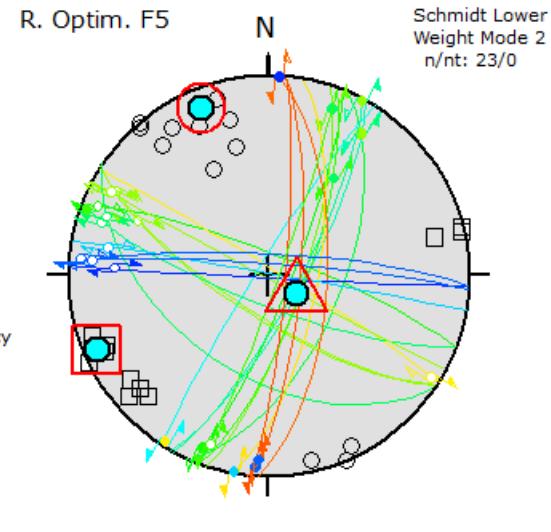
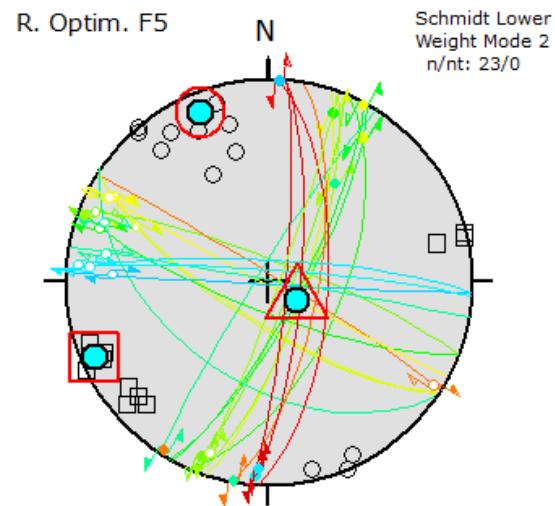


# Se identifican y separan 2 estados de esfuerzos coexistentes Con diferentes calidades y magnitudes no absolutas

## STRIKE-SLIP al NW

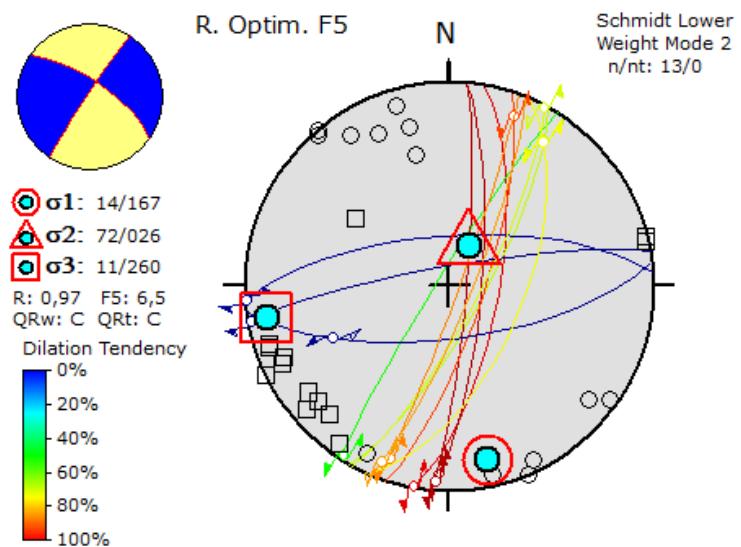
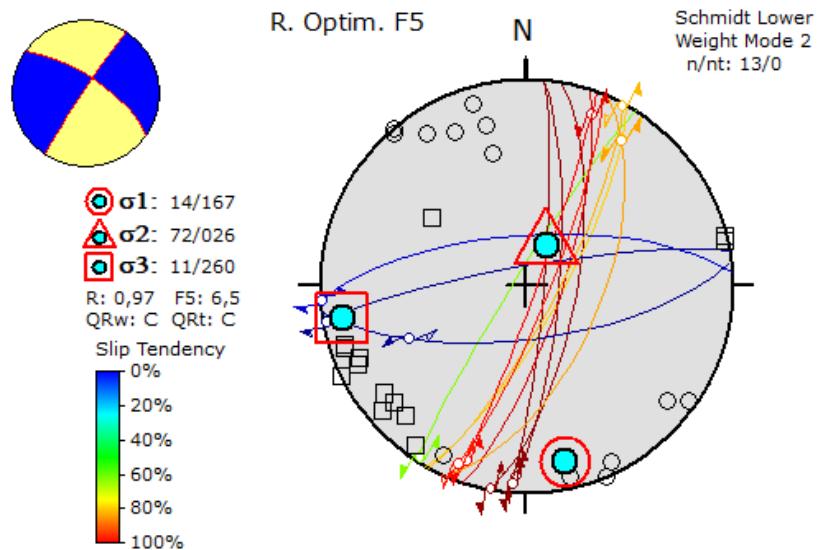
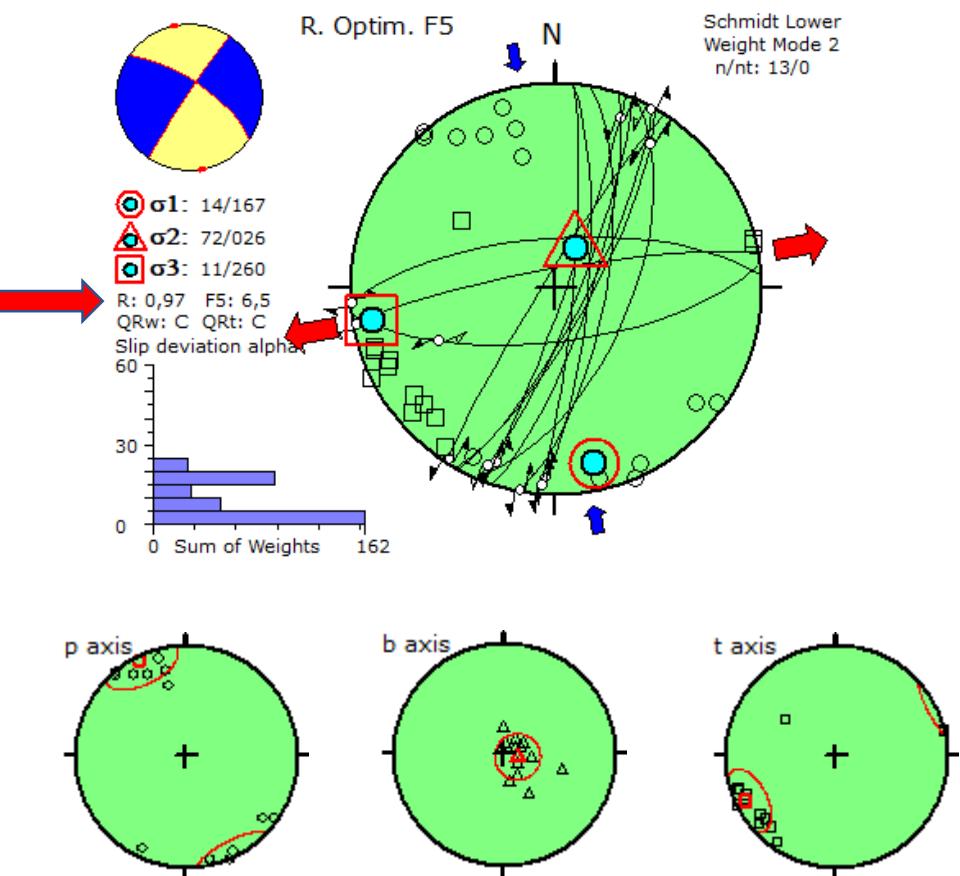


- $\sigma_1: 11/338$
  - △  $\sigma_2: 77/122$
  - $\sigma_3: 08/247$
- R: 0,65 F5: 5,1  
QRfm: A
- Slip Tendency  
0%  
20%  
40%  
60%  
80%  
100%



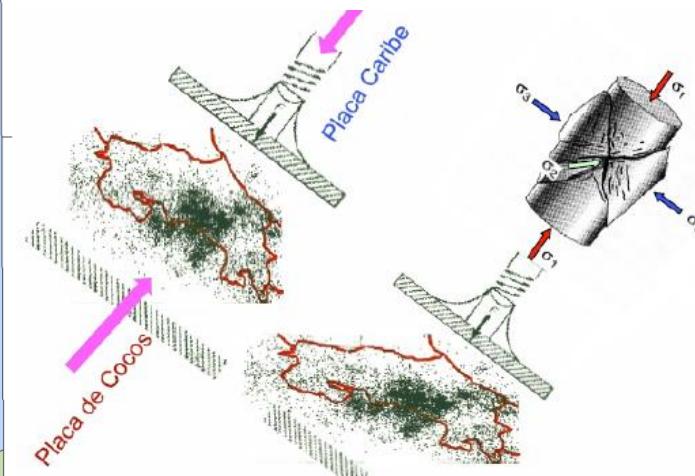
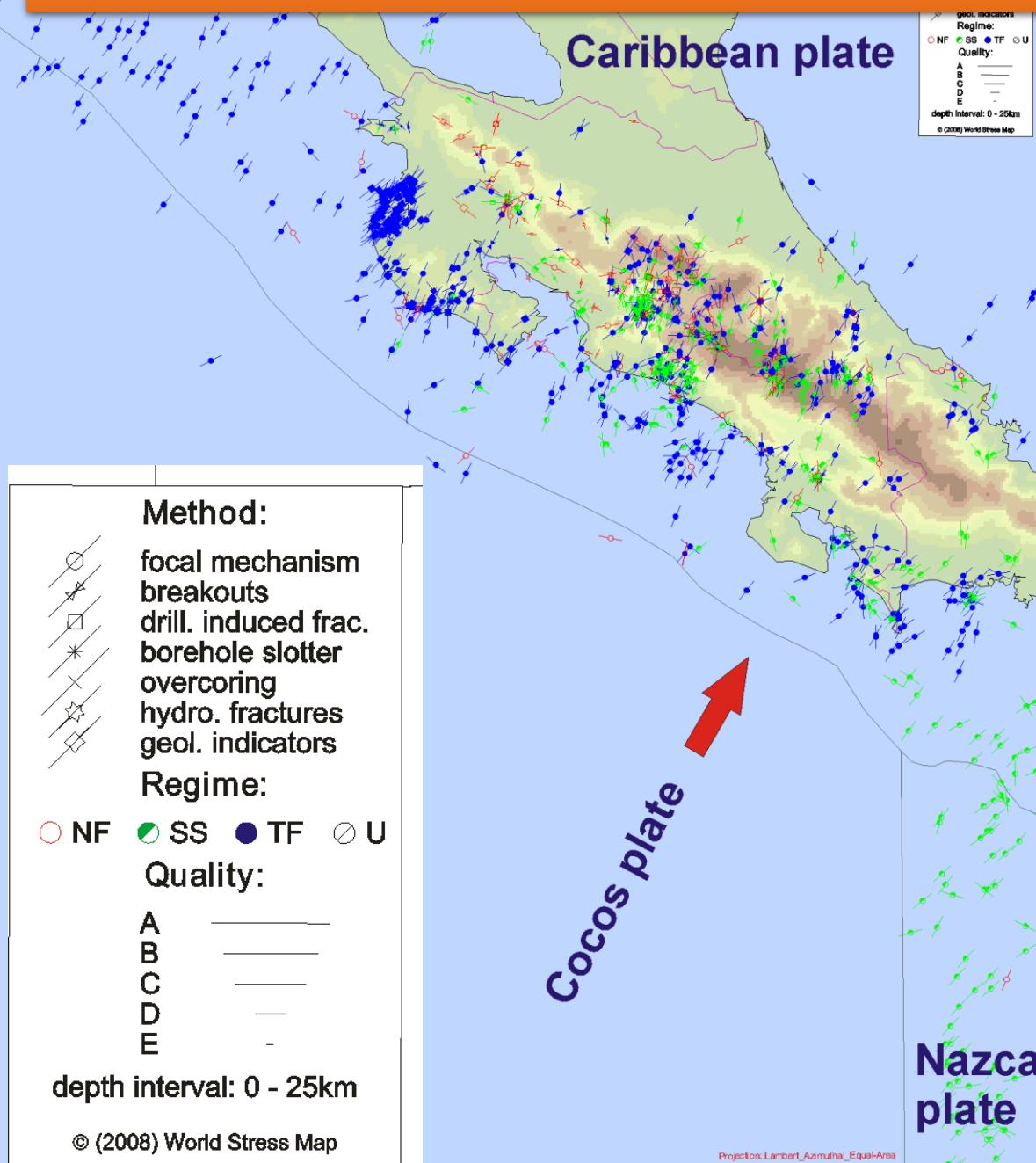
$R' = 1.03$

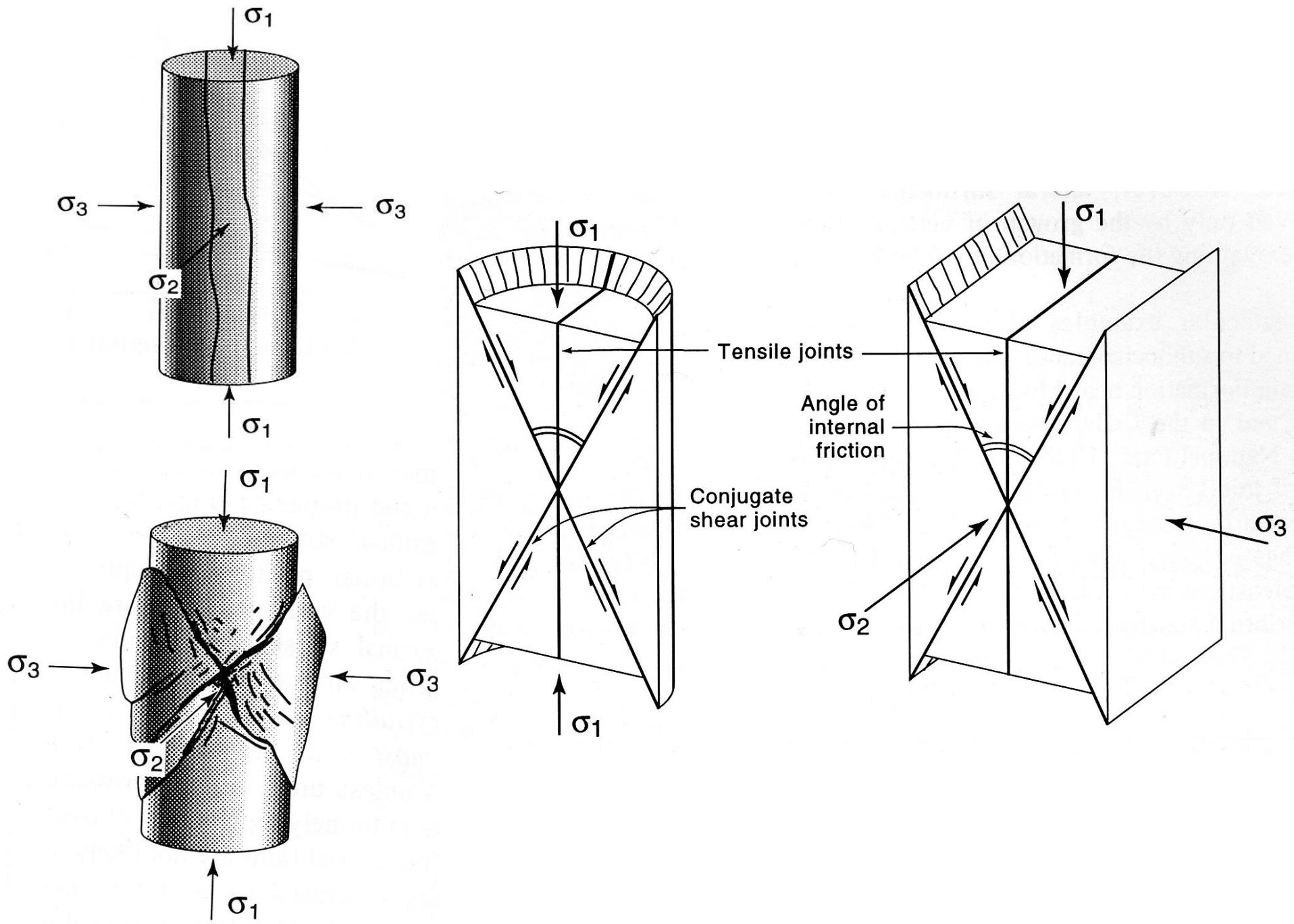
## STRIKE-SLIP al NWW, menor calidad © y muy inestable

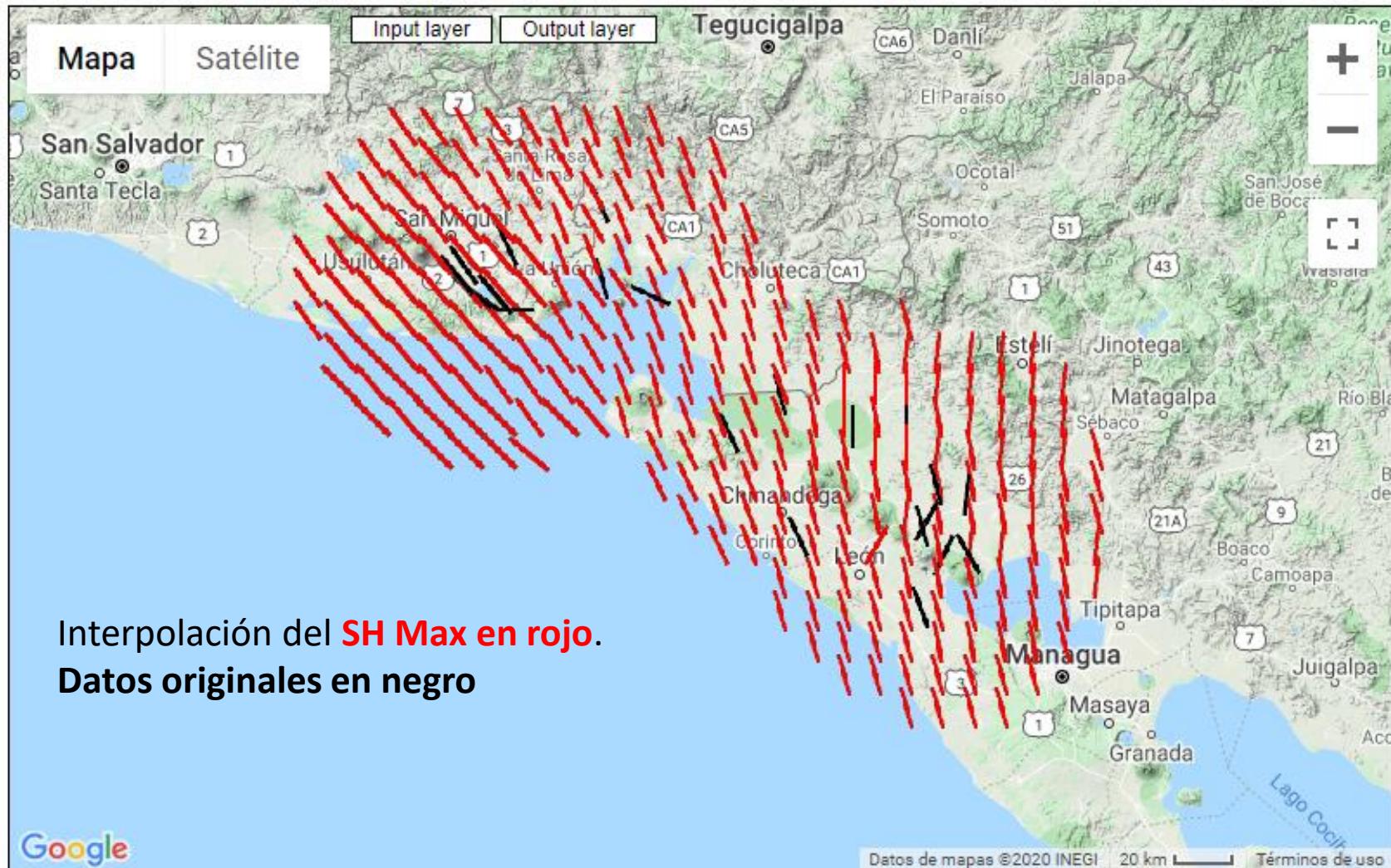


Esperaba un régimen compresivo  
con fallas-mecanismos inversos  
asociados a la subducción  
( hasta los 25 Km de profundidad)

Se investigará el estado con  
hipocentros a mayores  
profundidades







<http://shine.rm.ingv.it/index.phtml>



1. Data selection:  
Custom file

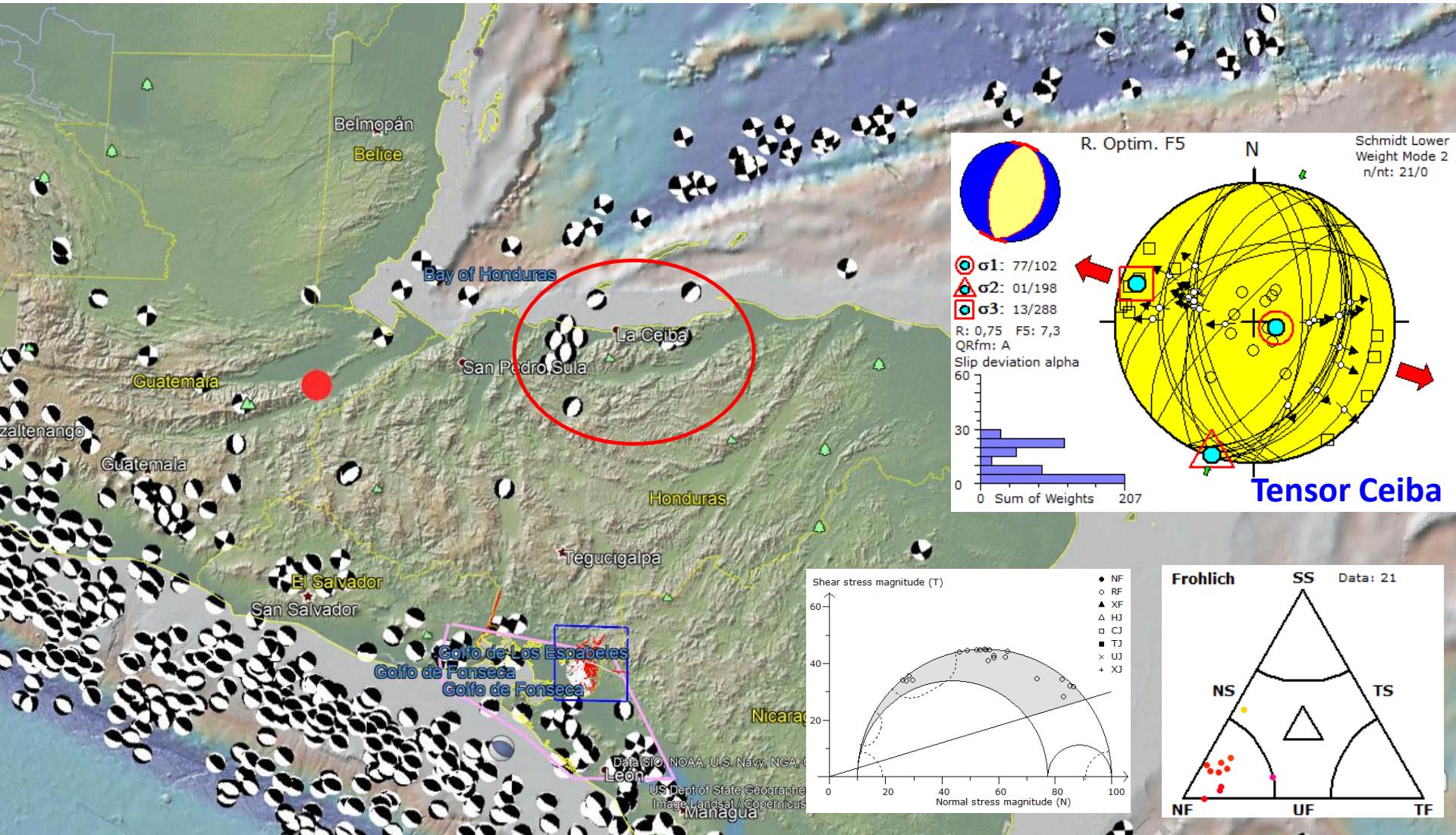


2. Geographical setting:  
Longitude: -89.43°/-85.78°  
Latitude: 11.95°/14.36°  
Grid spacing: 0.1°

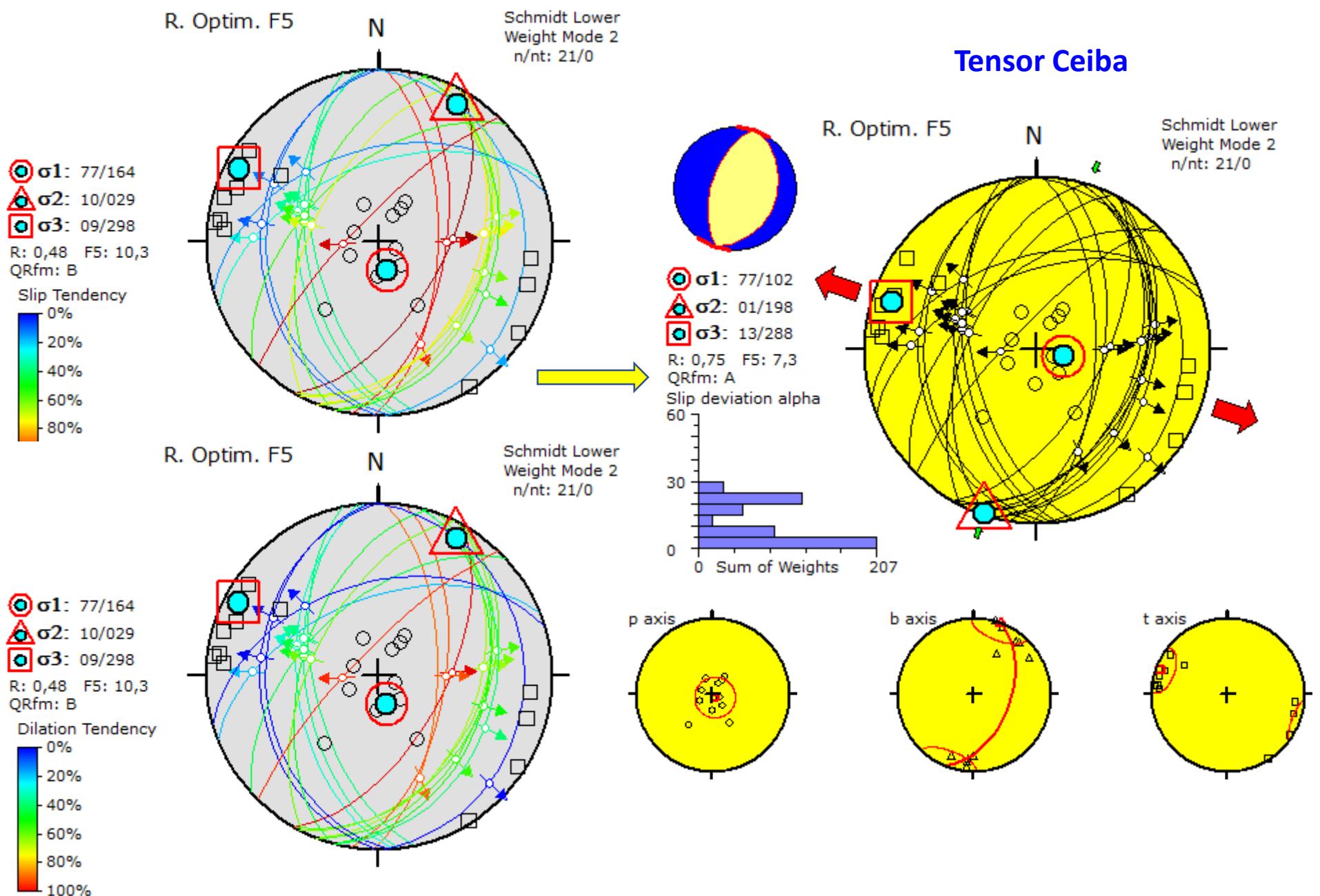


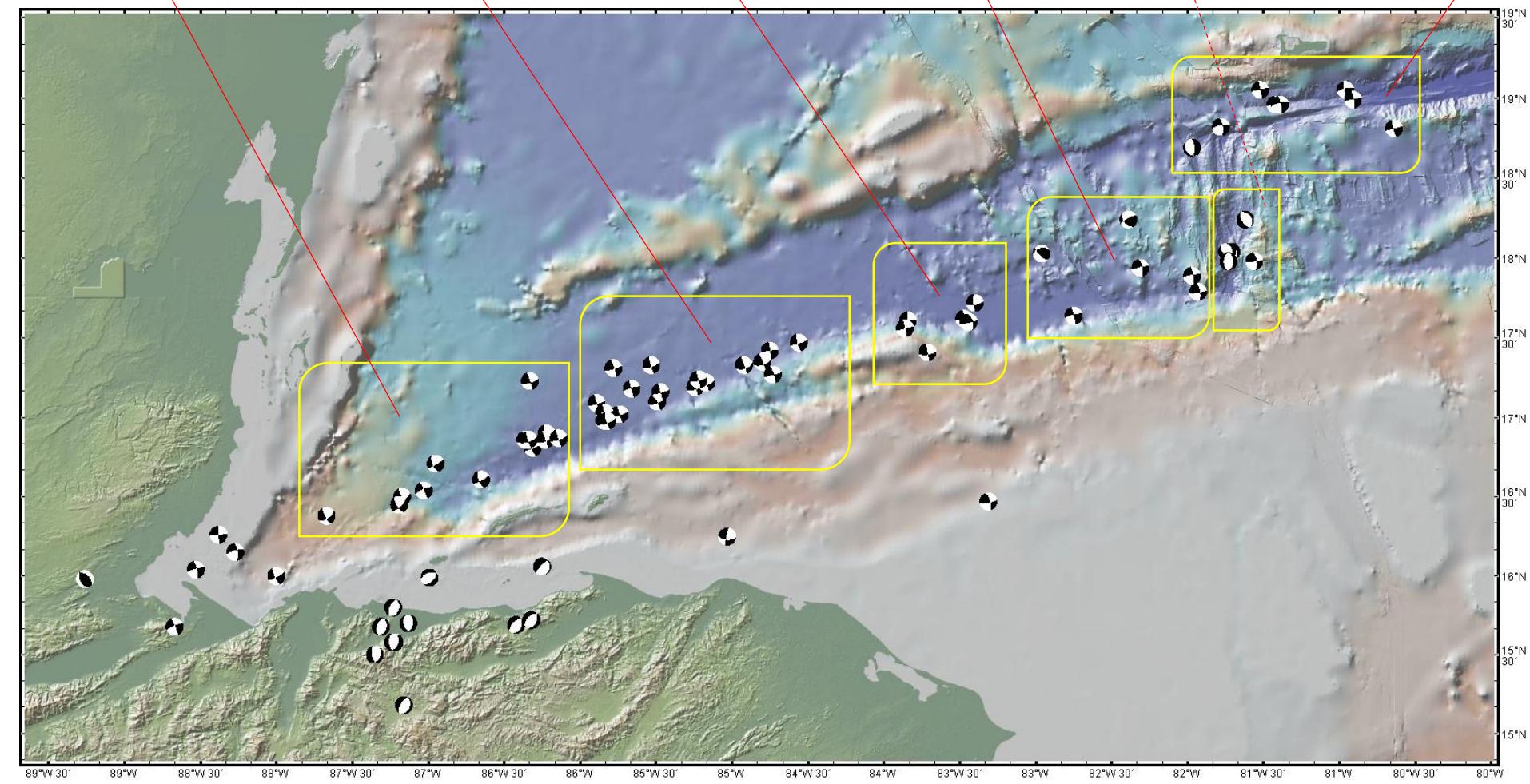
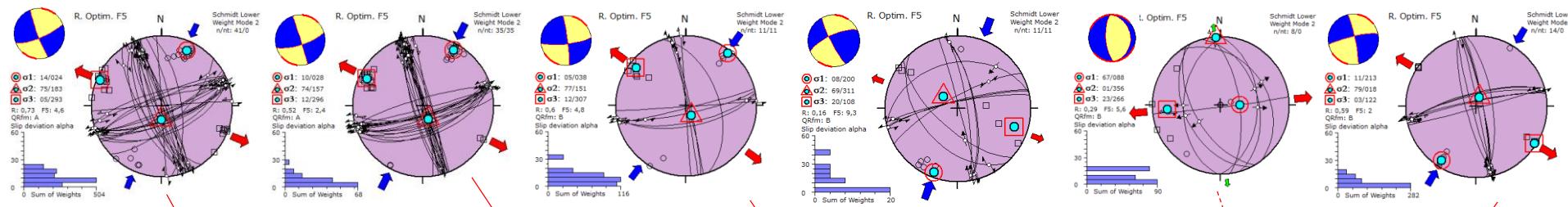
3. Strategic parameters:  
Searching radius: 2  
Minimum cluster: 3  
90% confidence bounds: 40°

## 2 Planos nodales de cada evento Mw > 4.0

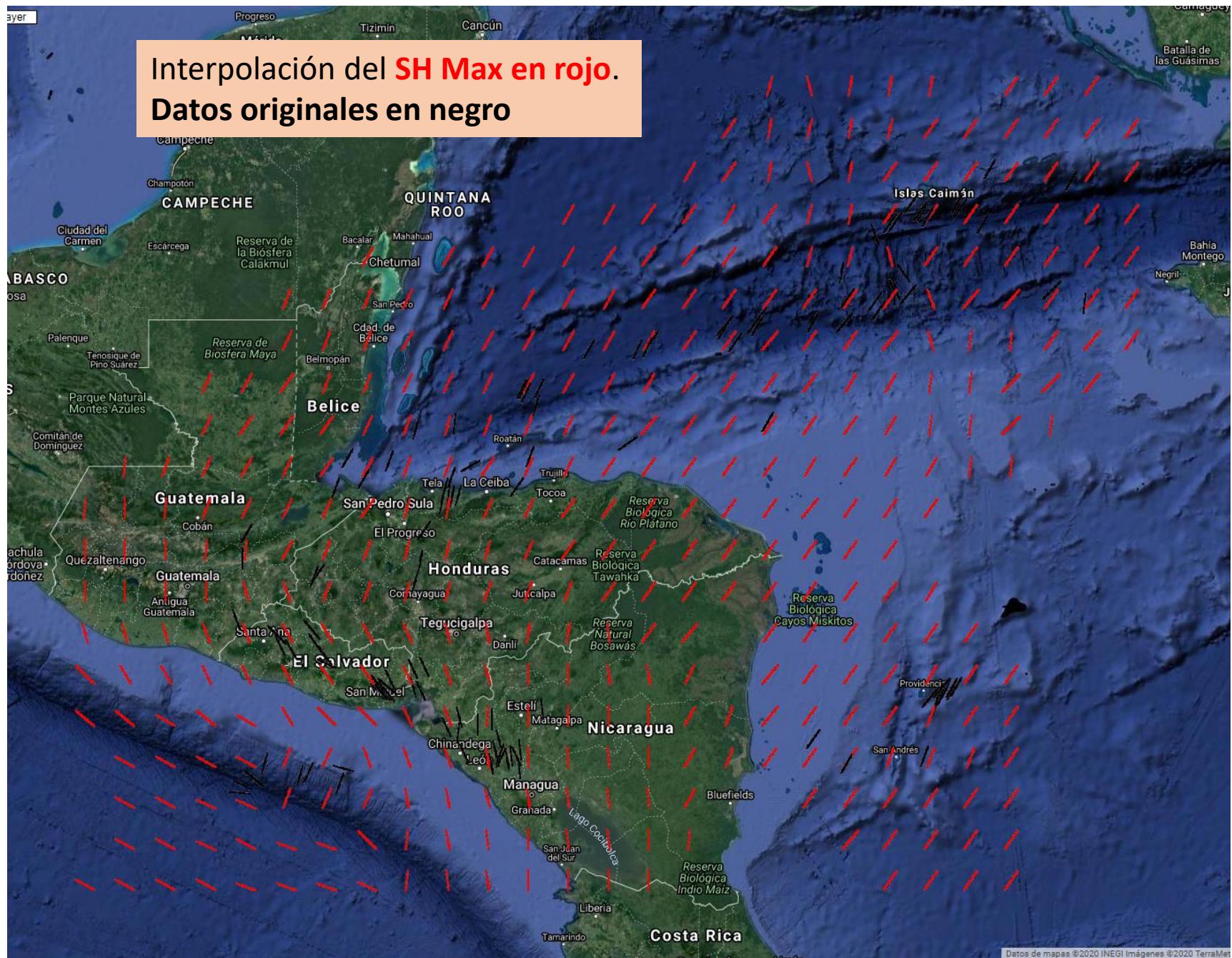


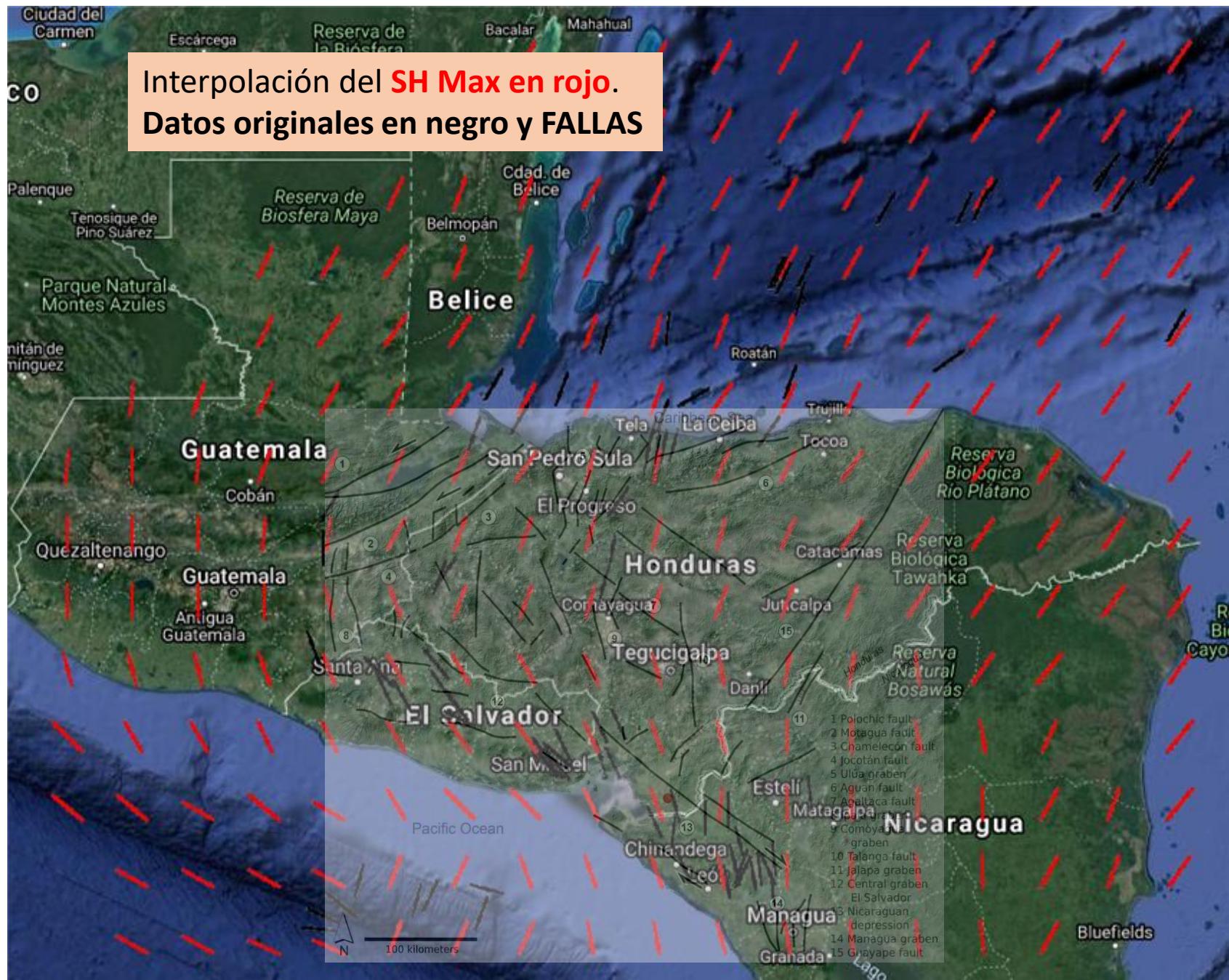
## 2 Planos nodales de cada evento Mw > 4.0





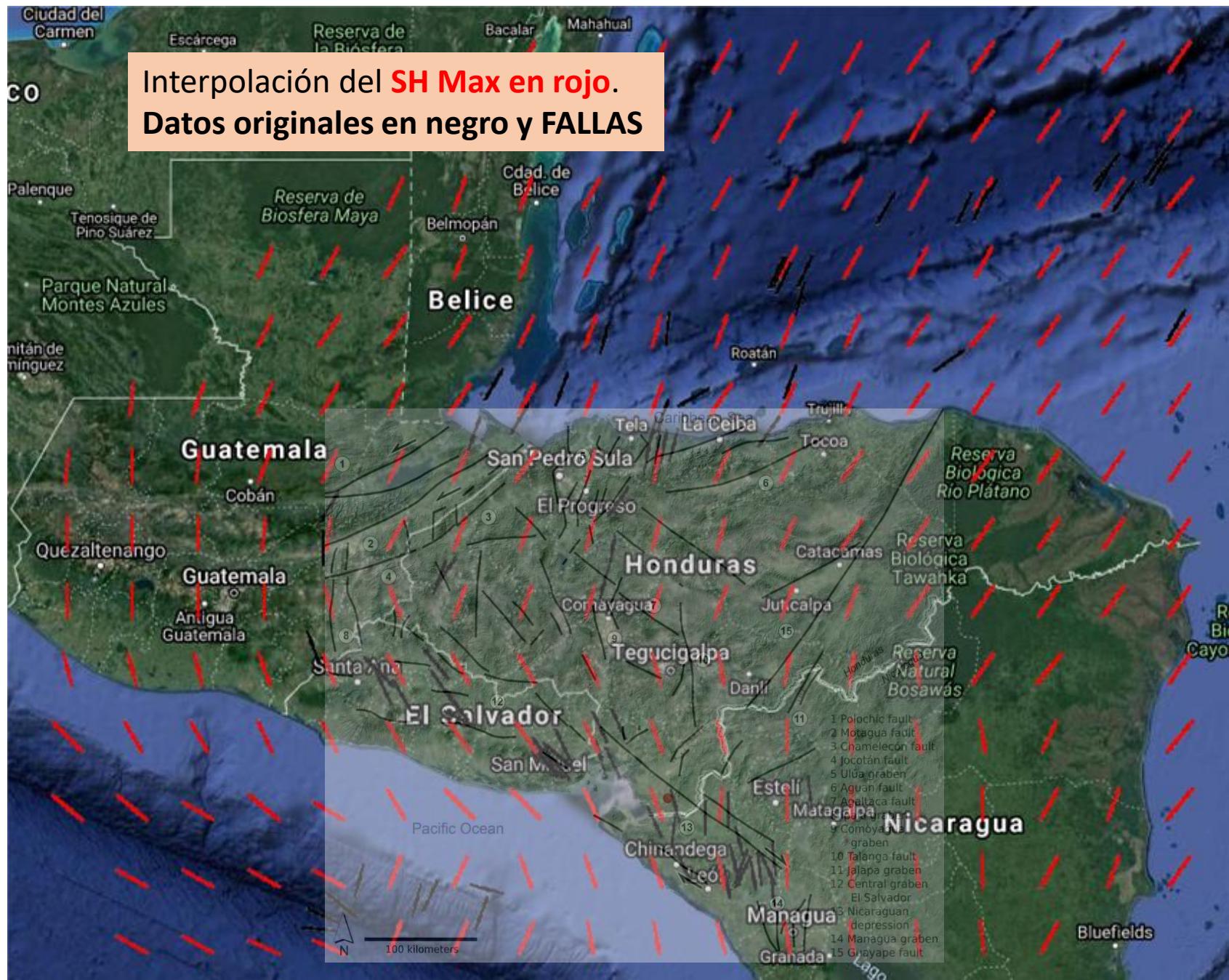
# Interpolación del SH Max en rojo. Datos originales en negro





# Interpolación del SH Max en rojo. Datos originales en negro





# Aplicaciones

# Definitions for some reservoir petrophysical properties derived from the acting stress tensor



Earth and Planetary Science Letters

Volume 478, 15 November 2017, Pages 159-174



Igneous sills record far-field and near-field  
stress interactions during volcano  
construction: Isle of Mull, Scotland

T.L. Stephens <sup>a</sup>  , R.J. Walker <sup>a</sup>, D. Healy <sup>b</sup>, A. Bubeck <sup>a</sup>, R.W. England <sup>a</sup>, K.J.W. McCaffrey <sup>c</sup>

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<https://doi.org/10.1016/j.epsl.2017.09.003>

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**Slip and dilation tendency** show the attitude of planes susceptible to reactivation, via shear or dilation respectively.

A plane is susceptible to slip when the ratio of shear stress ( $\tau$ ) to the normal stress acting on the plane ( $\sigma_n$ ) is large. Slip tendency is normalised here (relative to the maximum possible slip tendency,  $T_s(\max)$ ) to enable comparison between stress states (Eq. (1); Morris et al., 1996):

$$(1) \quad T_s = (\tau / \sigma_n) / T_s(\max)$$

A plane is susceptible to dilation when the difference between  $\sigma_1$  and the normal stress acting on the plane approaches the magnitude of differential stress ( $\sigma_D$ , where  $\sigma_D = \sigma_1 - \sigma_3$ ; Eq. (2); Ferrill et al., 1999):

$$T_d = \frac{(\sigma_1 - \sigma_r)}{(\sigma_1 - \sigma_3)}$$

**Fracture susceptibility** (Eq. (3); e.g. [Mildren et al., 2002](#)) represents the magnitude of fluid pressure change ( $\Delta P_f$ : either magmatic pressure within a crack, or hydrous pore-fluid pressure within a crack or the host rock pore space) that is required to cause shear reactivation.

Fluid-driven shear reactivation is dependent on the shear and normal stresses acting on the plane, as well as its cohesion (here assumed = 0), and static coefficient of friction ( $\mu_s$ ):

$$(3) \quad S_f = \sigma n - (\tau / \mu_s)$$

Reactivation via fluid-driven dilation would require a fluid pressure greater than the fracture susceptibility.

Intrusion is favoured where  $T_d$  is high and fracture susceptibility is low.

Shear reactivation is favoured where  $T_s$  is high and fracture susceptibility is low.

The ***opening angle*** is related to the

***shear stress ( $\tau$ ),  
normal stress ( $\sigma_n$ ), and  
fluid pressure ( $P_f$ )***

$$\mu = \tan^{-1} \left( \frac{\tau}{P_f - \sigma_n} \right)$$

acting on that plane at the time of intrusion (Delaney et al., 1986; Jolly and Sanderson, 1997):

Equation (5) shows that if the fluid overpressure ( $P_f - \sigma_n$ ) is equal to the shear stress, the opening angle is  $45^\circ$ , and the shear-to-dilation ratio is unity. If the overpressure is greater than the shear stress, the opening angle is less than  $45^\circ$ , and the fracture will show a greater component of dilation to shear.

***When  $\mu$  is negative, the fracture will remain closed as the fluid pressure did not exceed the normal stress.***

An intrusive segment, however, may inflate against a closed fracture (where  $P_f < \sigma_n$ ), causing a local contractional shear and a blunt intrusion tip.

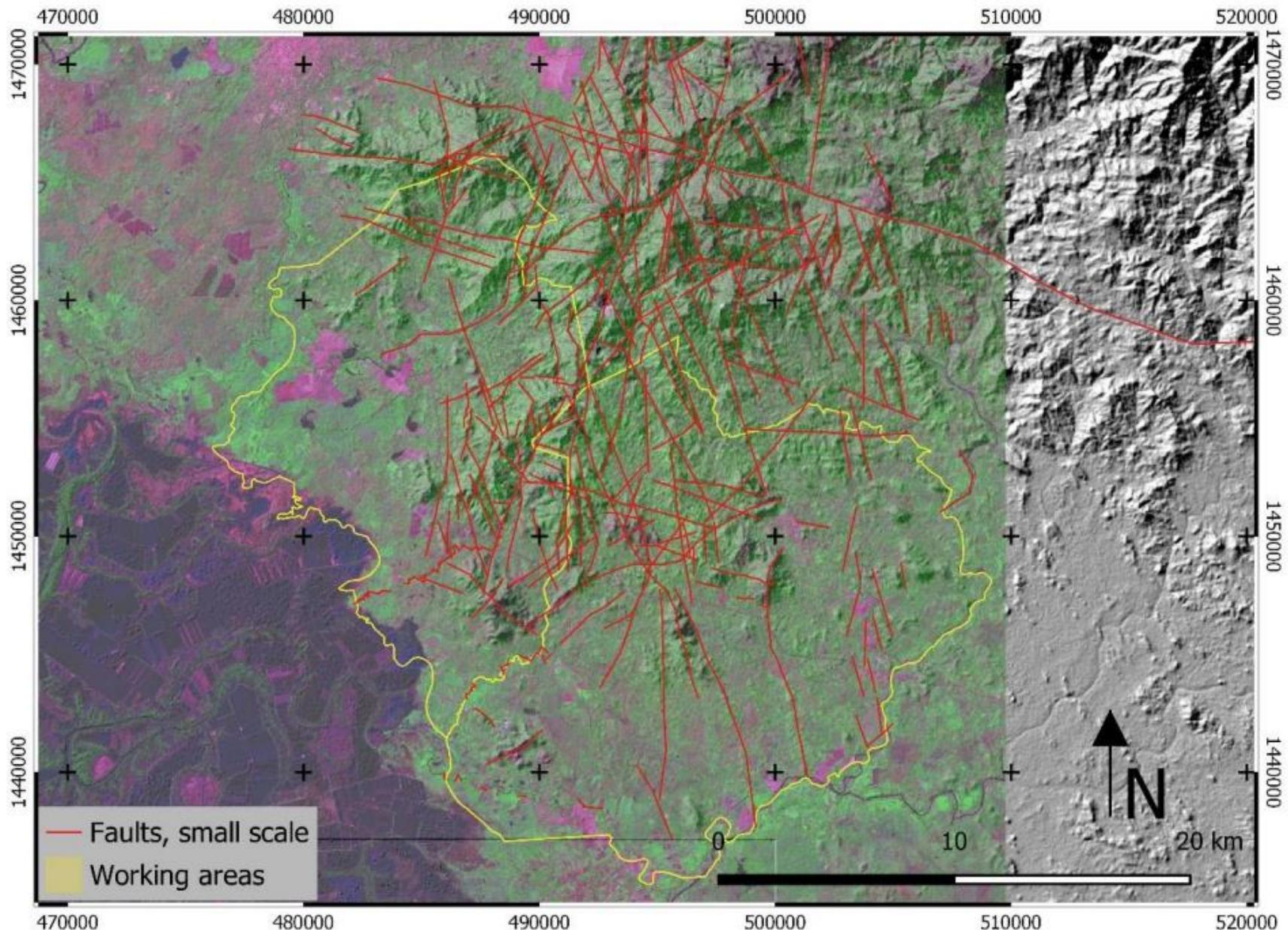
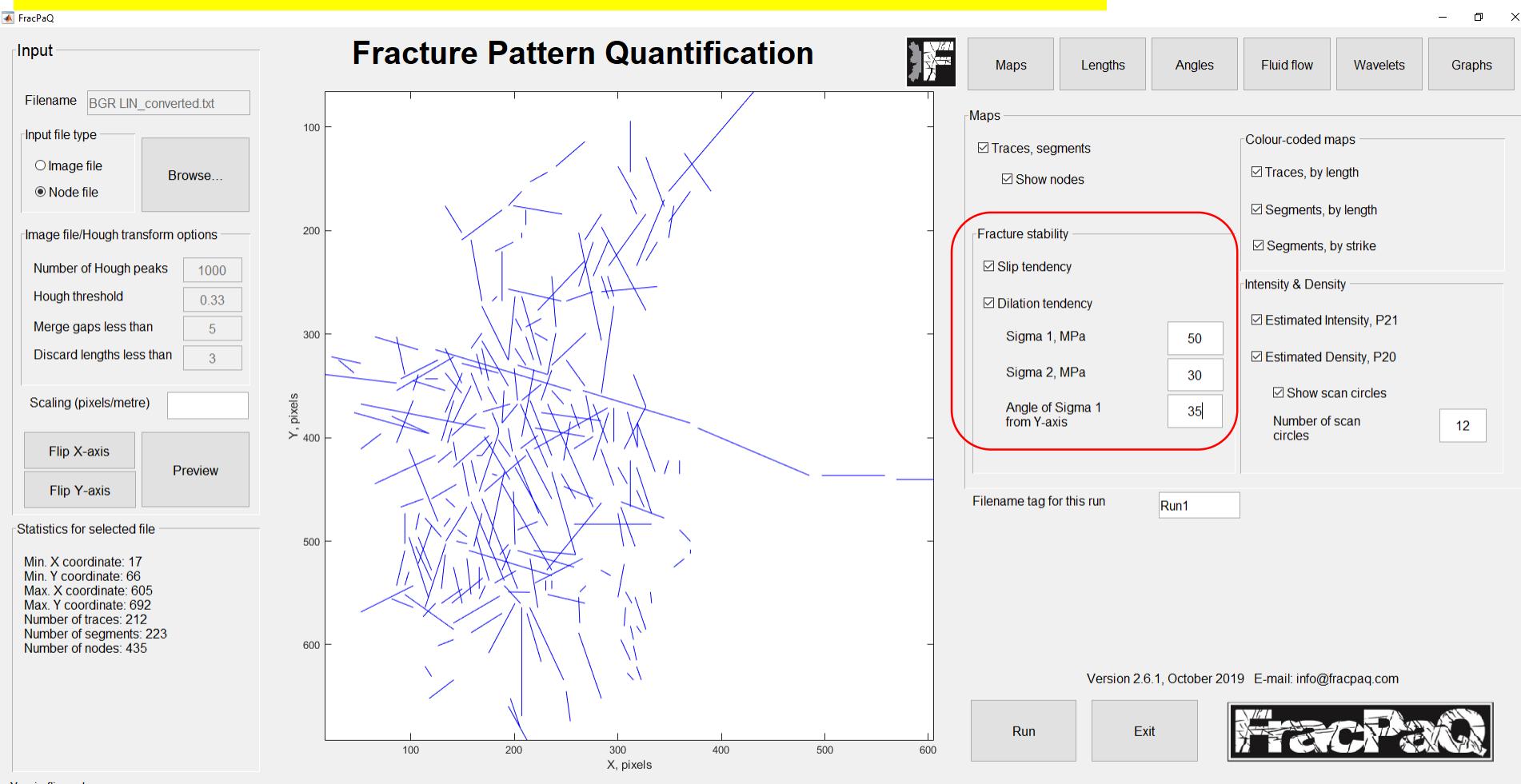


Figure 5.4: Lineaments interpreted as faults of the working areas and the surrounding areas in a small scale. Sentinel 2A, bands 12, 8a, 5 (RGB) over shaded relief SRTM DEM.

# Lineamientos identificados por BGR, 2019 en Namasigüe-El Triunfo (Alina Ermertz & Dr. Kai Hahne )



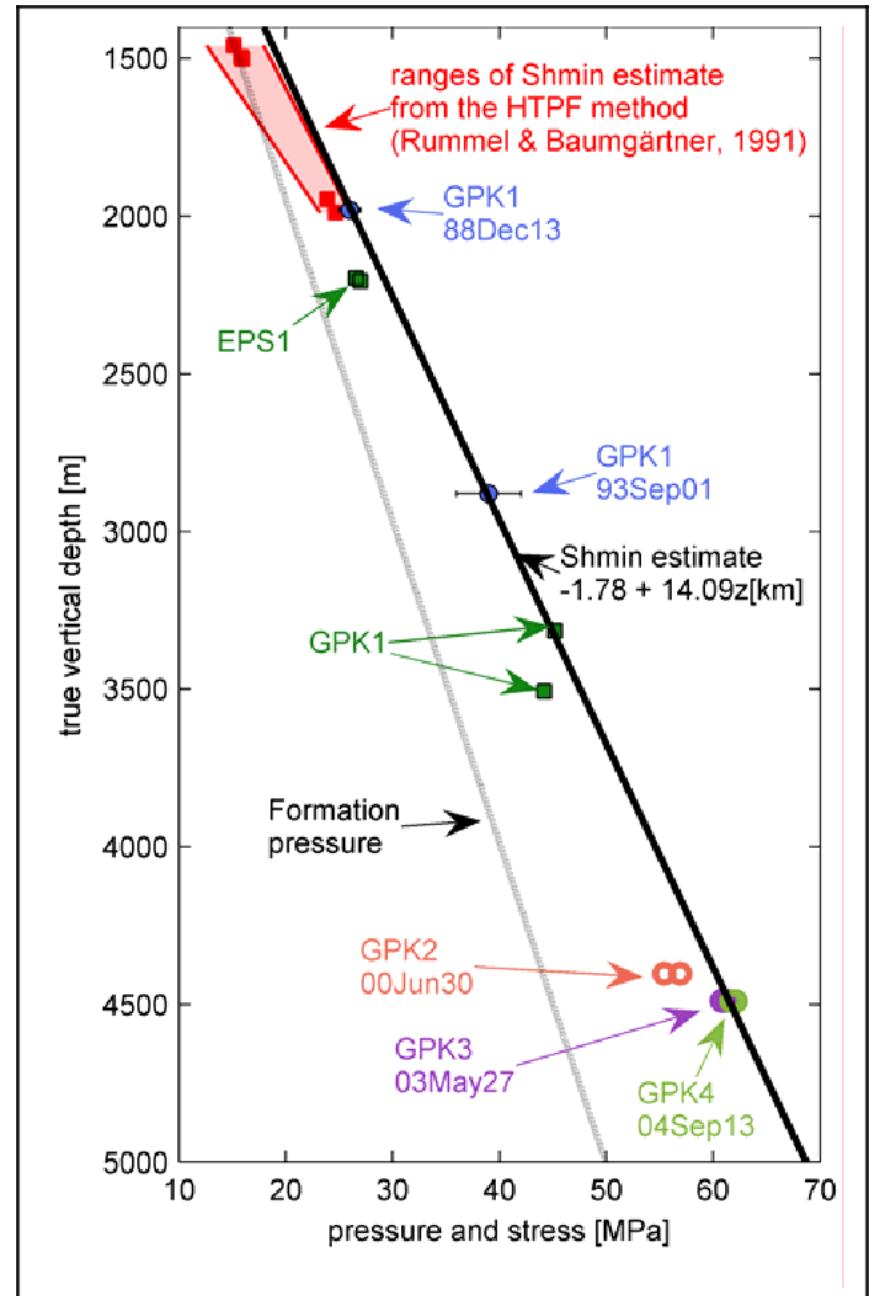
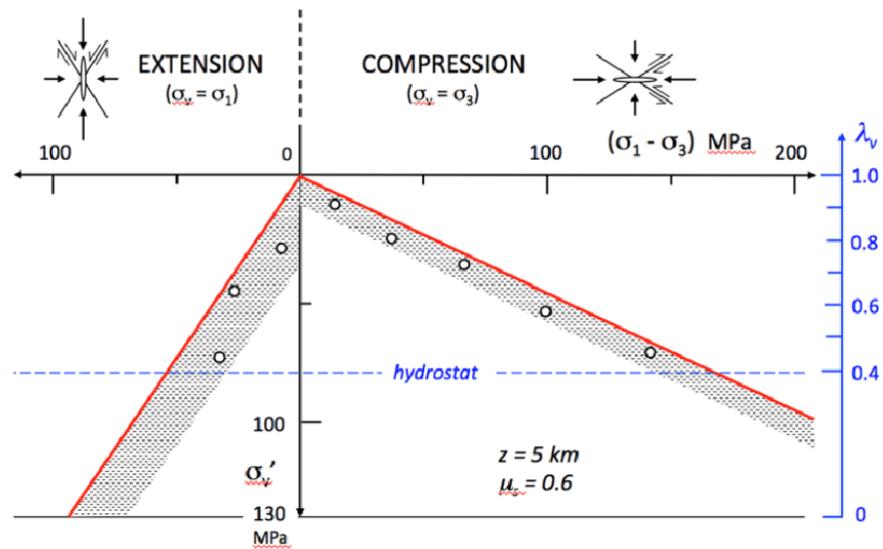
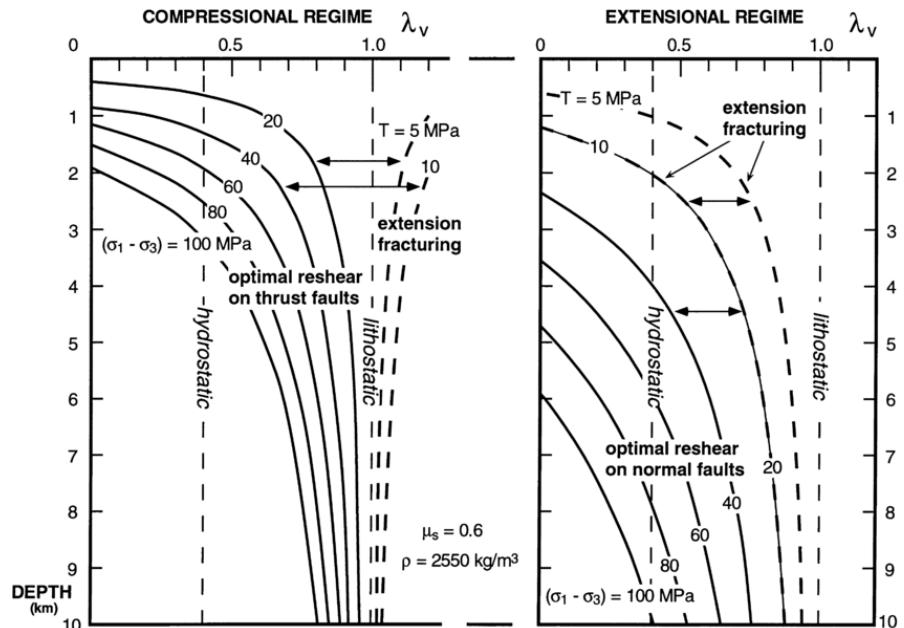
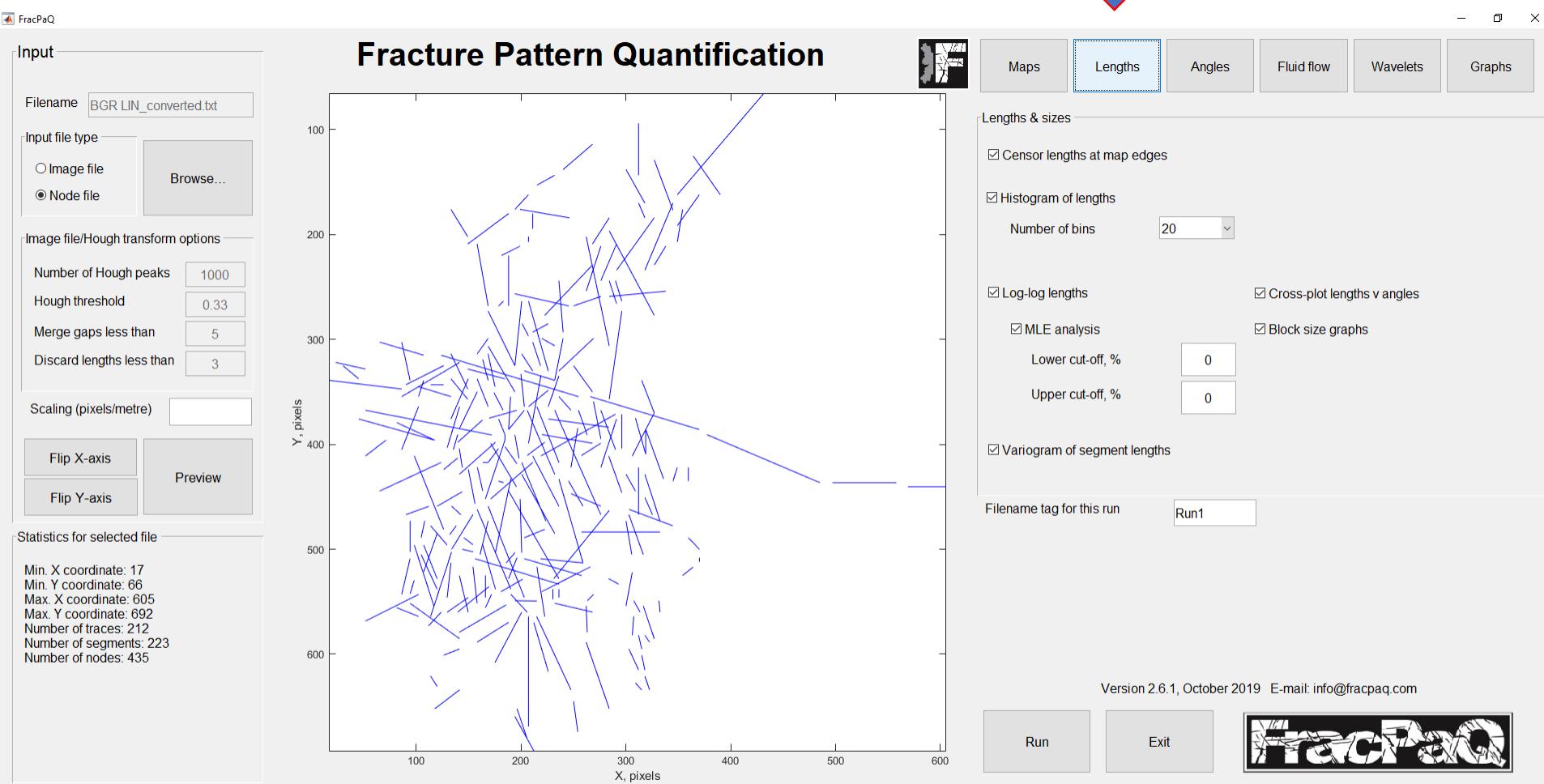
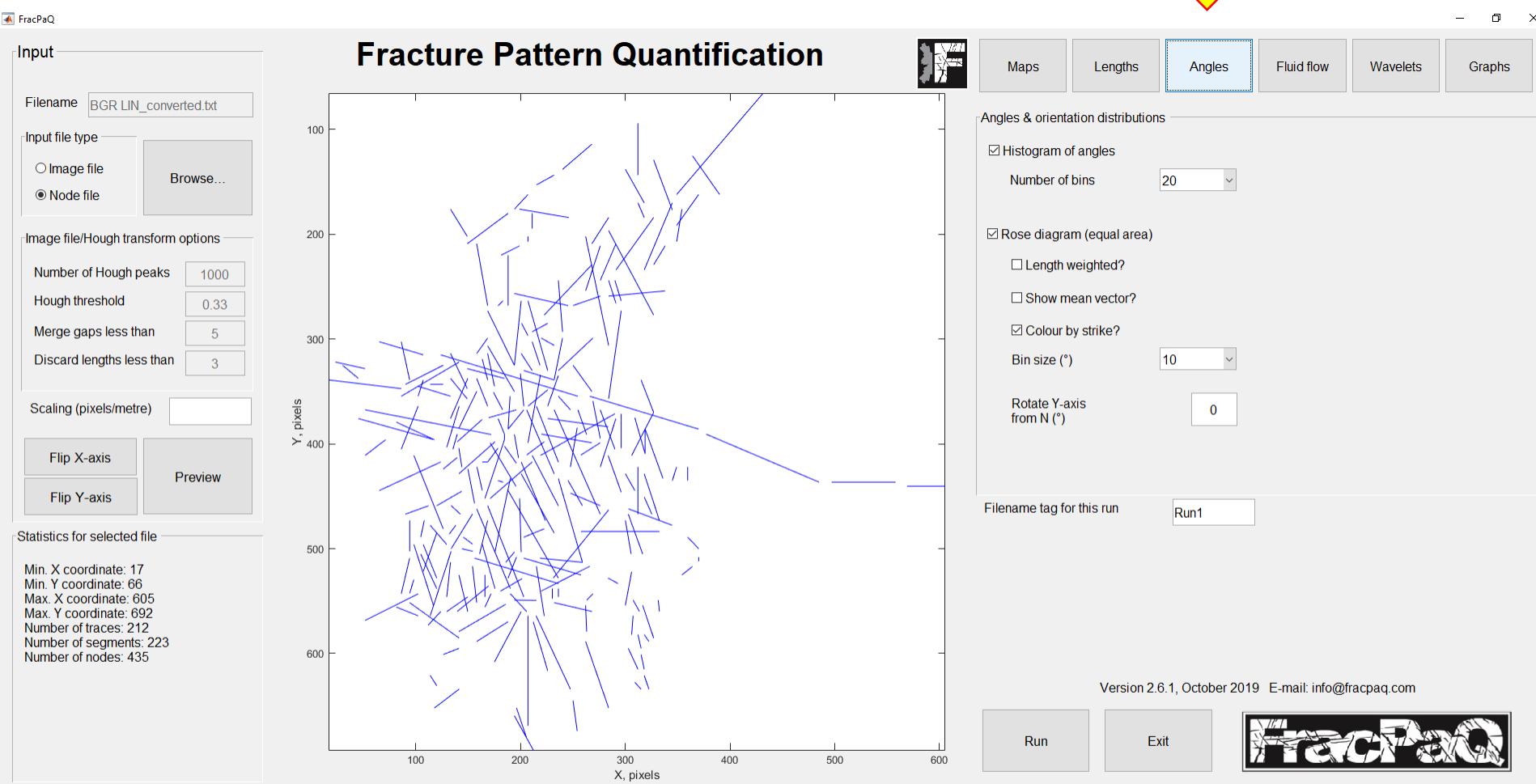


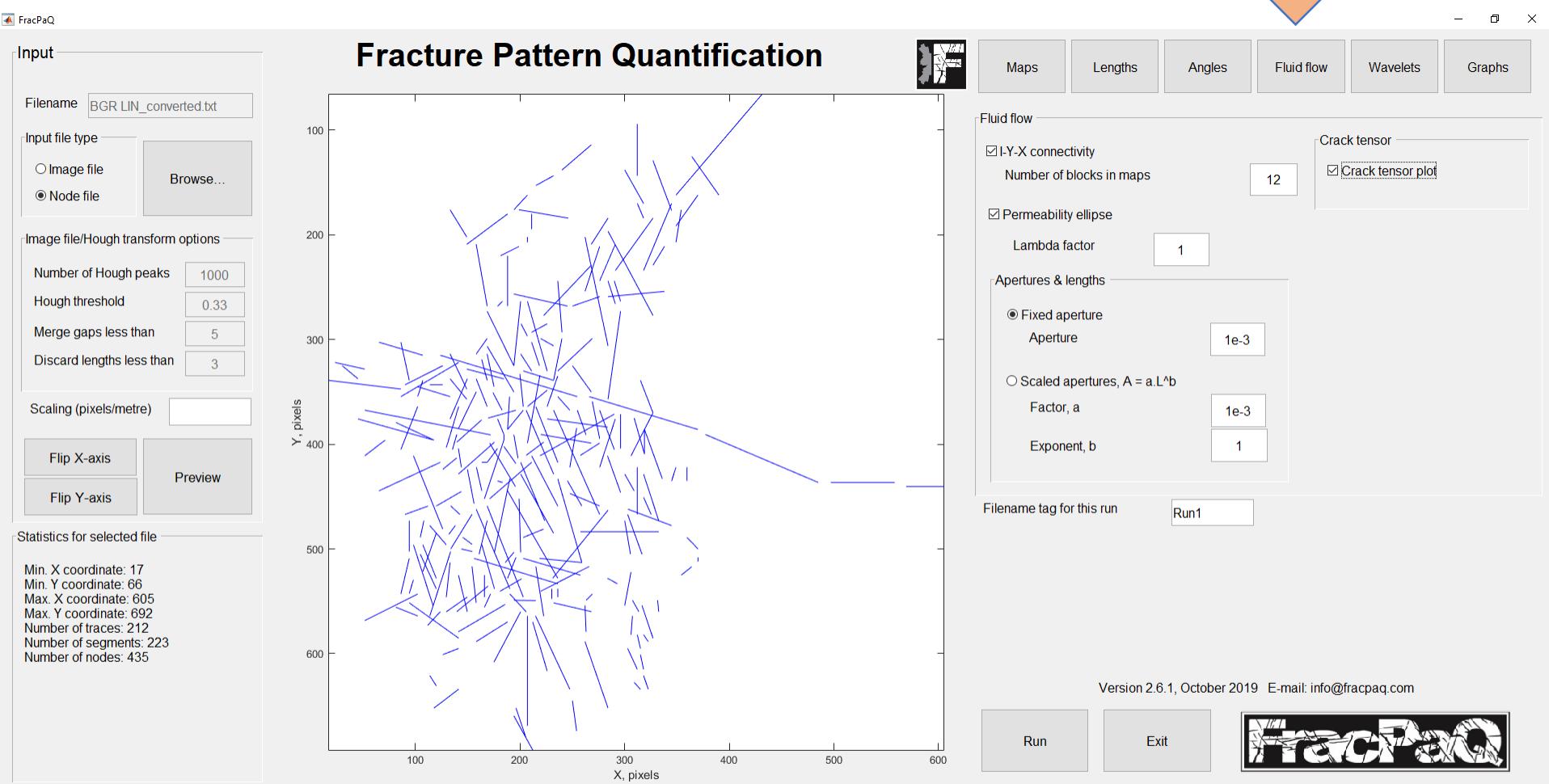
Figure 4: Estimates of the Shmin from small-volume

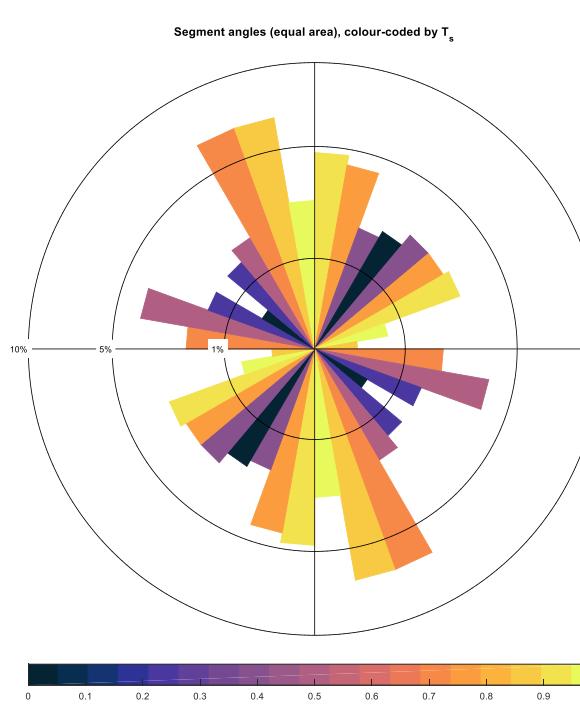
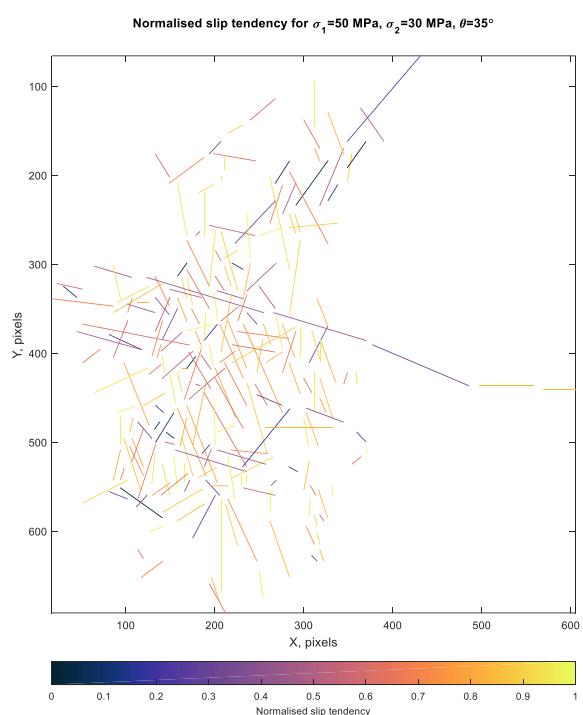




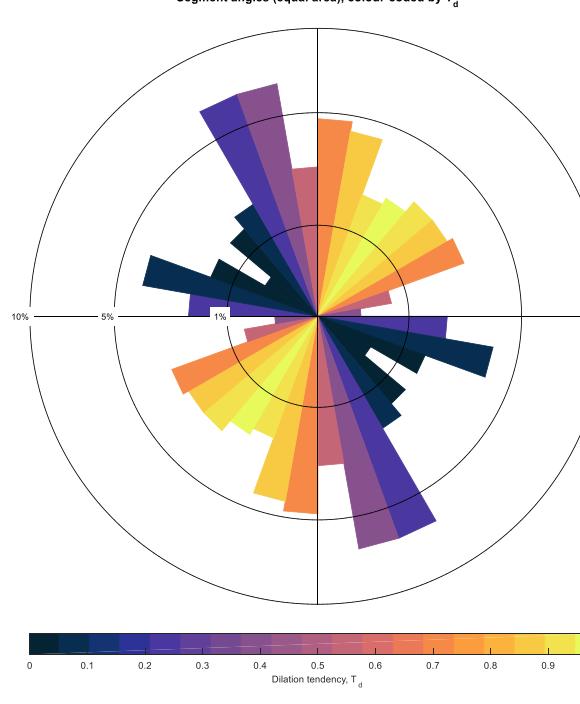
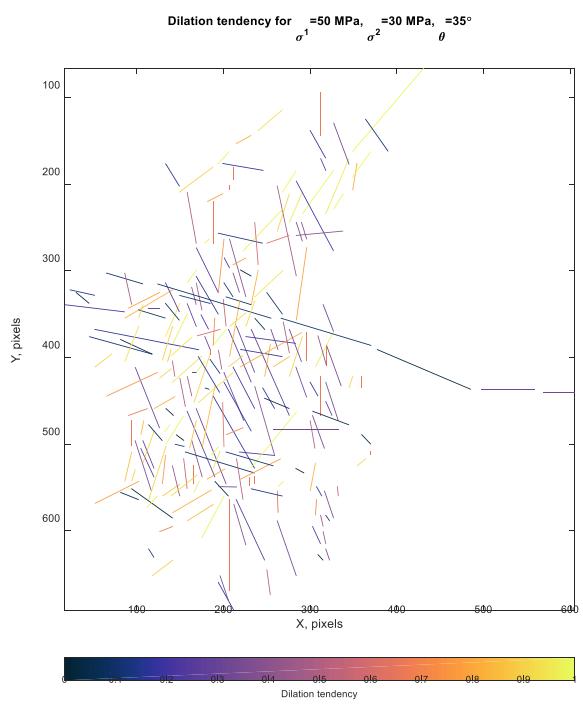
Version 2.6.1, October 2019 E-mail: info@fracpaq.com





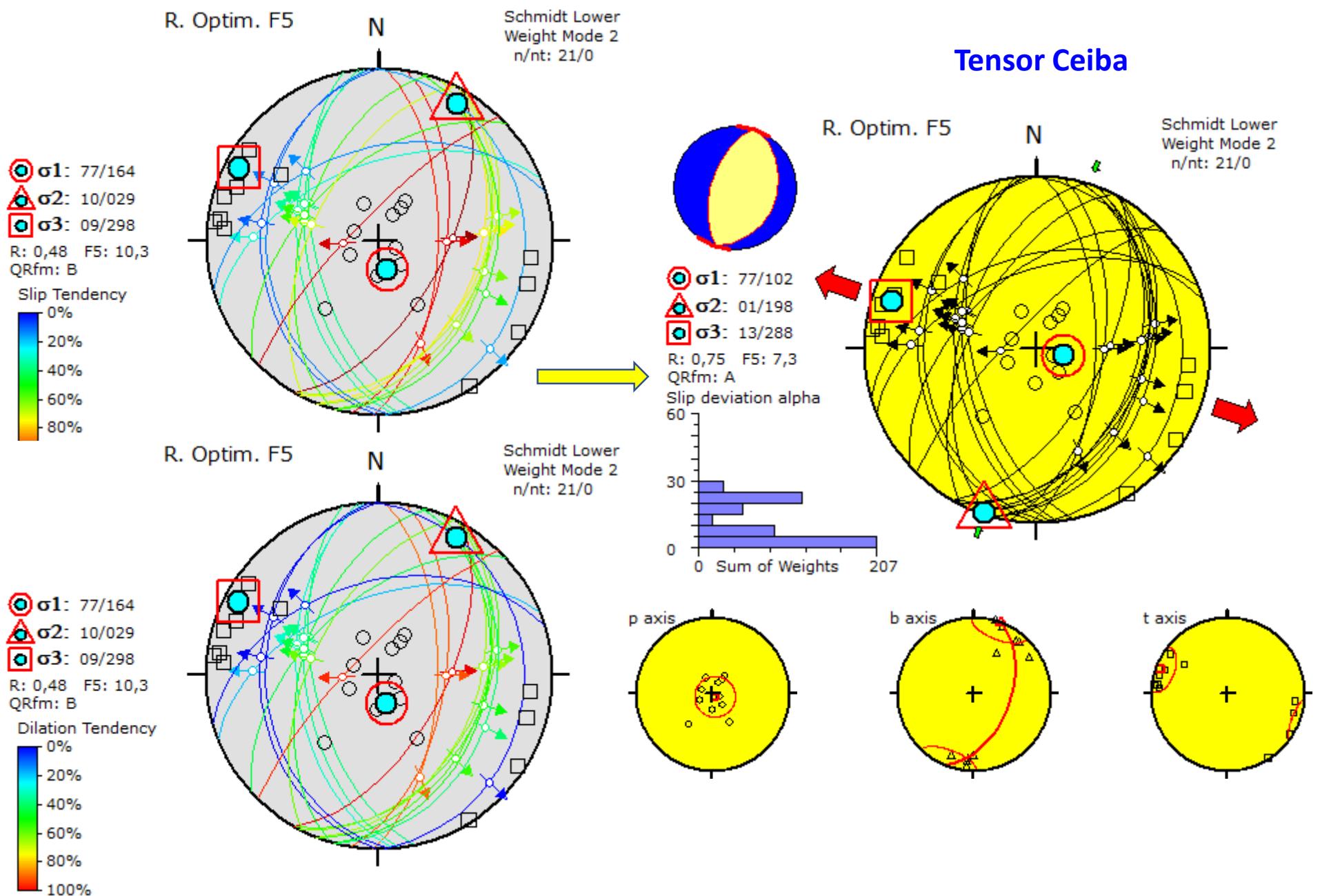


$$T_s = (\tau/\sigma n)/T_s(\max)$$



$$T_d = \frac{(\sigma_1 - \sigma_r)}{(\sigma_1 - \sigma_3)}$$

## 2 Planos nodales de cada evento Mw > 4.0



Muchas gracias  
por su paciencia

*Preguntas ?*

*Dudas ?*

*Sospechas ?*